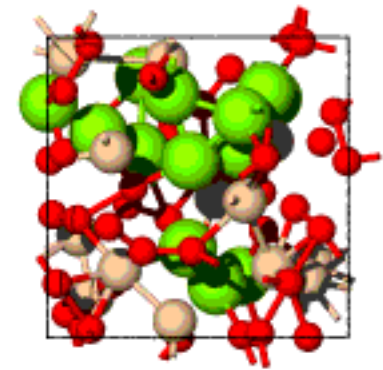


Quantum simulations at extreme conditions: warm dense matter and planetary interiors

Presentation to:
LLNL HEDS Seminar Series

Felipe González

Department of Earth and Planetary Science
University of California, Berkeley



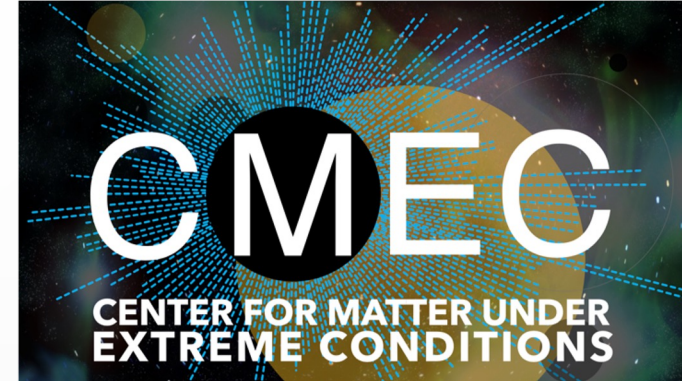
Berkeley
UNIVERSITY OF CALIFORNIA

12/01/2022

ACKNOWLEDGEMENTS



DOE-NNSA
(DE-NA0003842)



Collaborators:

B. Militzer (UCB)

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K. Driver (LLNL)

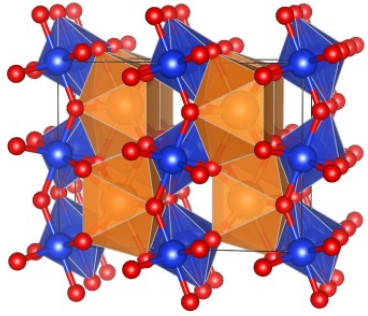
R. Jeanloz (UCB)

B.K. Godwal (UCB)



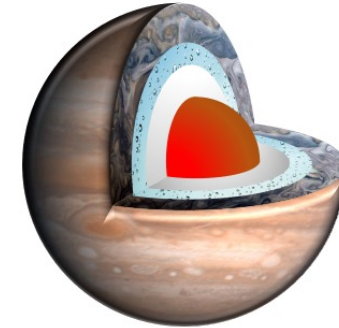
National Energy Research
Scientific Computing Center

OVERVIEW AND MOTIVATION



High pressure

- Solid-solid phase transitions
- Equations of state
- Melting

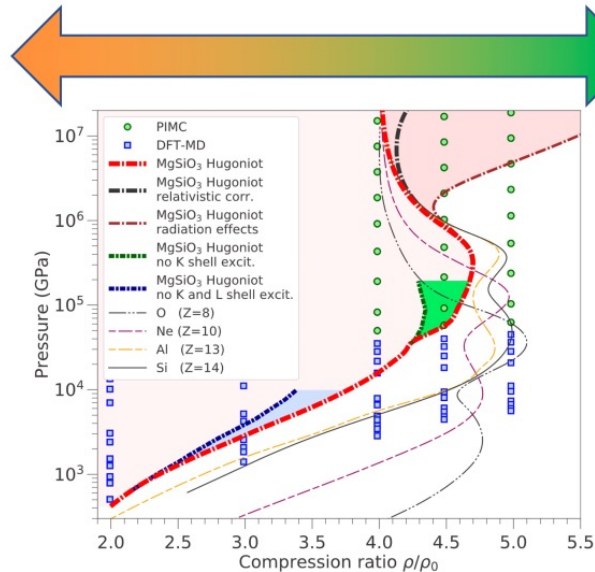


Warm dense matter

- Temperature ionization
- Pressure ionization
- Shock compression

Planetary Interiors

- Super-Earth interior models
- Core/mantle crystallization
- Element Partitioning
- Solubility / Miscibility



OUTLINE

1. Planetary Interiors

- A diluted core in Jupiter
- Rock/Ice mixtures in water planets

2. High Pressure Phase Transitions

- Be & MgO: melting and anharmonicities
- Melting of SiO_2
- Ramp compression from DFT

3. Warm Dense Matter




- Warm dense silicates: Mg, MgO & MgSiO_3
- FPEOS




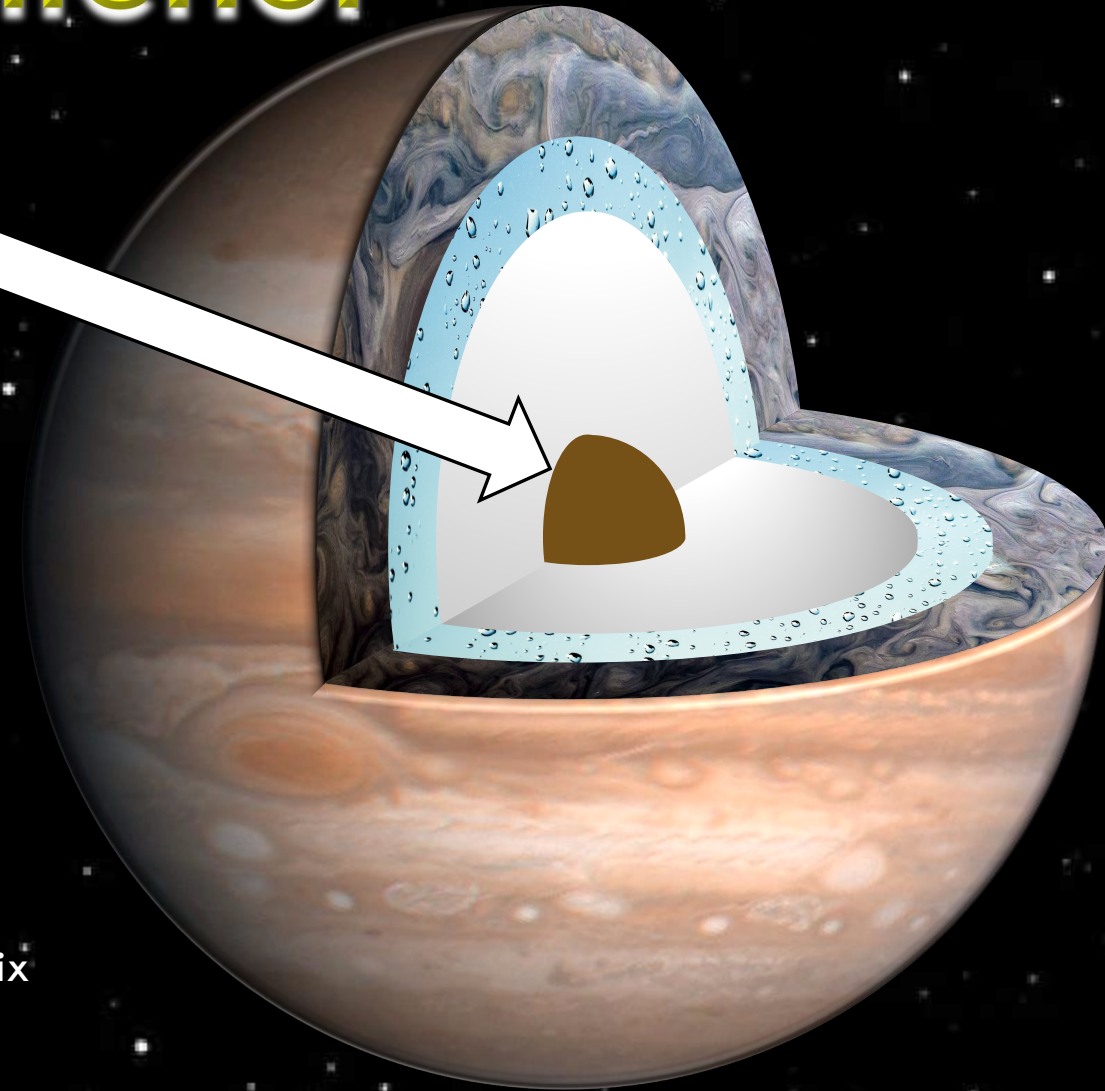
Jupiter's Interior

$P \sim 40 \text{ Mbar}$ (4000 GPa)
 $T \sim 16000 \text{ K}$

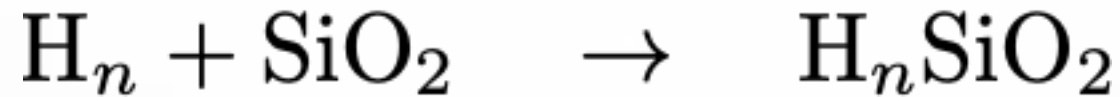
Can metallic H
dissolve the rocky core?

 H + He mix
 He rain
 Metallic H

 Compact core
(Solid/Liquid?)

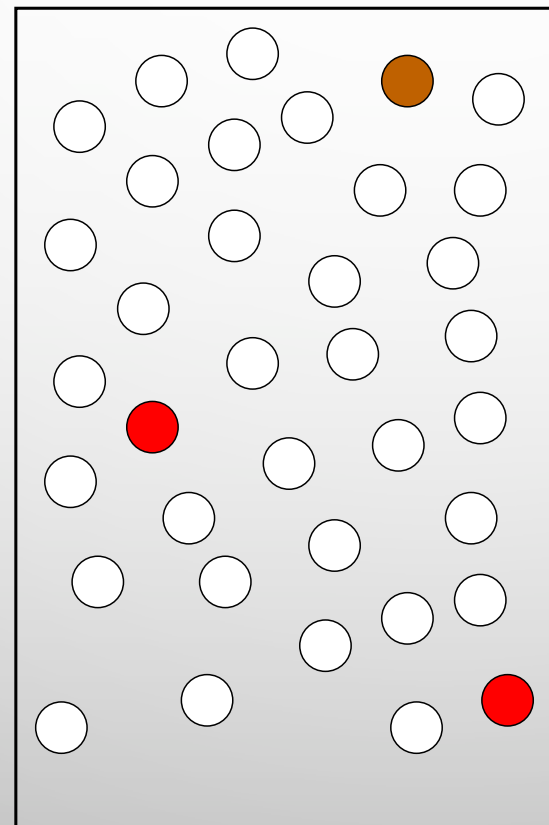
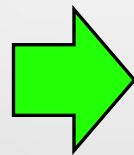
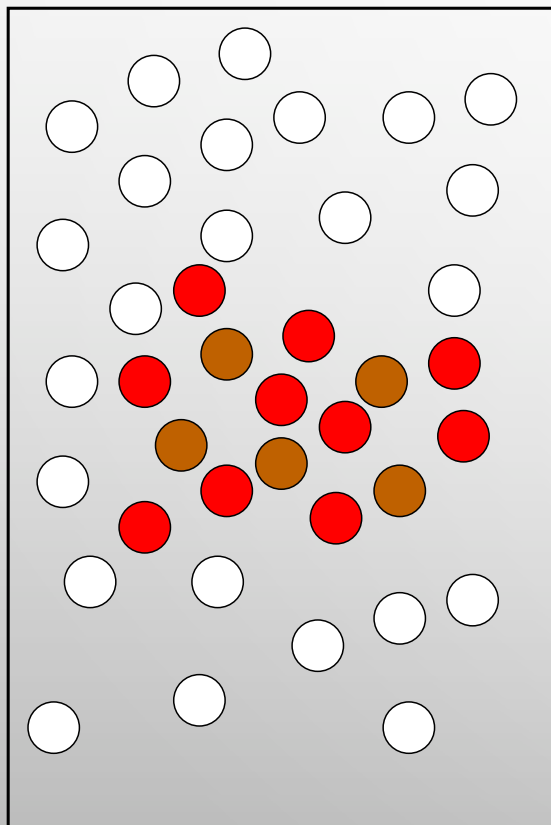


Solubility of SiO_2



Free energy of solvation

$$\Delta G \equiv G(\text{H}_n\text{SiO}_2) - [G(\text{H}_n) + G(\text{SiO}_2)]$$



Free energy of the
Dissolved System

$$\Delta G < 0 \Rightarrow \text{Dissolves}$$

How to calculate $G(P,T)$?

Thermodynamic Integration

$$U_A \rightarrow U_B$$

U_A : reference system

F_A : known

$$U(\lambda) = U_A + \lambda(U_B - U_A)$$

$$(U(0) = U_A; U(1) = U_B)$$

U_B : DFT system

F_B : unknown

$$\Delta F = \int_A^B dF = \int_0^1 d\lambda \frac{\partial F}{\partial \lambda} = - \int_0^1 d\lambda \frac{k_B T}{Z} \frac{\partial Z}{\partial \lambda}$$

$$= \int_0^1 d\lambda \frac{k_B T}{Z} \sum_s e^{-\frac{U(\lambda)}{k_B T}} \frac{1}{k_B T} \frac{\partial U(\lambda)}{\partial \lambda}$$

$$= \int_0^1 d\lambda \left\langle \frac{\partial U(\lambda)}{\partial \lambda} \right\rangle_\lambda = \int_0^1 d\lambda \langle U_B - U_A \rangle_\lambda$$

$$F_B = F_A + \Delta F$$

$$G_B = F_B + PV$$

$$F = -kT \ln Z$$

$$Z = \sum_s e^{-\frac{U(\lambda)}{k_B T}}$$

PLANETARY INTERIORS

A diluted core in Jupiter

LIQUIDS



$$U = 0$$



TWO-STEPS TDI



$$U = U_{PP}$$



$$U = U_{KS}$$

- GGA-PBE
- Ecut = 900 eV
- 144 atoms
- Γ - point
- 14 val. elect.

SOLIDS



$$\frac{1}{2} m \omega r^2 \text{ (one body)}$$

$$U = U_{Einstein}$$

$\lambda = 0 \rightarrow \lambda = 1$



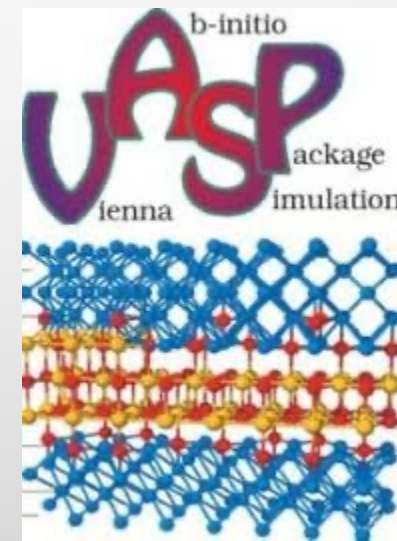
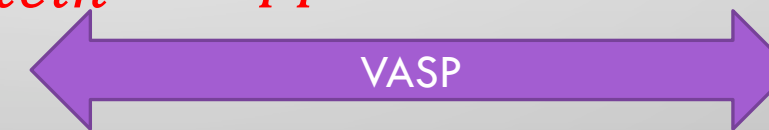
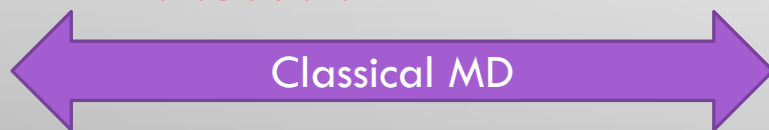
(one body + two body)

$$U = U_{Einstein} + U_{PP}$$

$\lambda = 0 \rightarrow \lambda = 1$

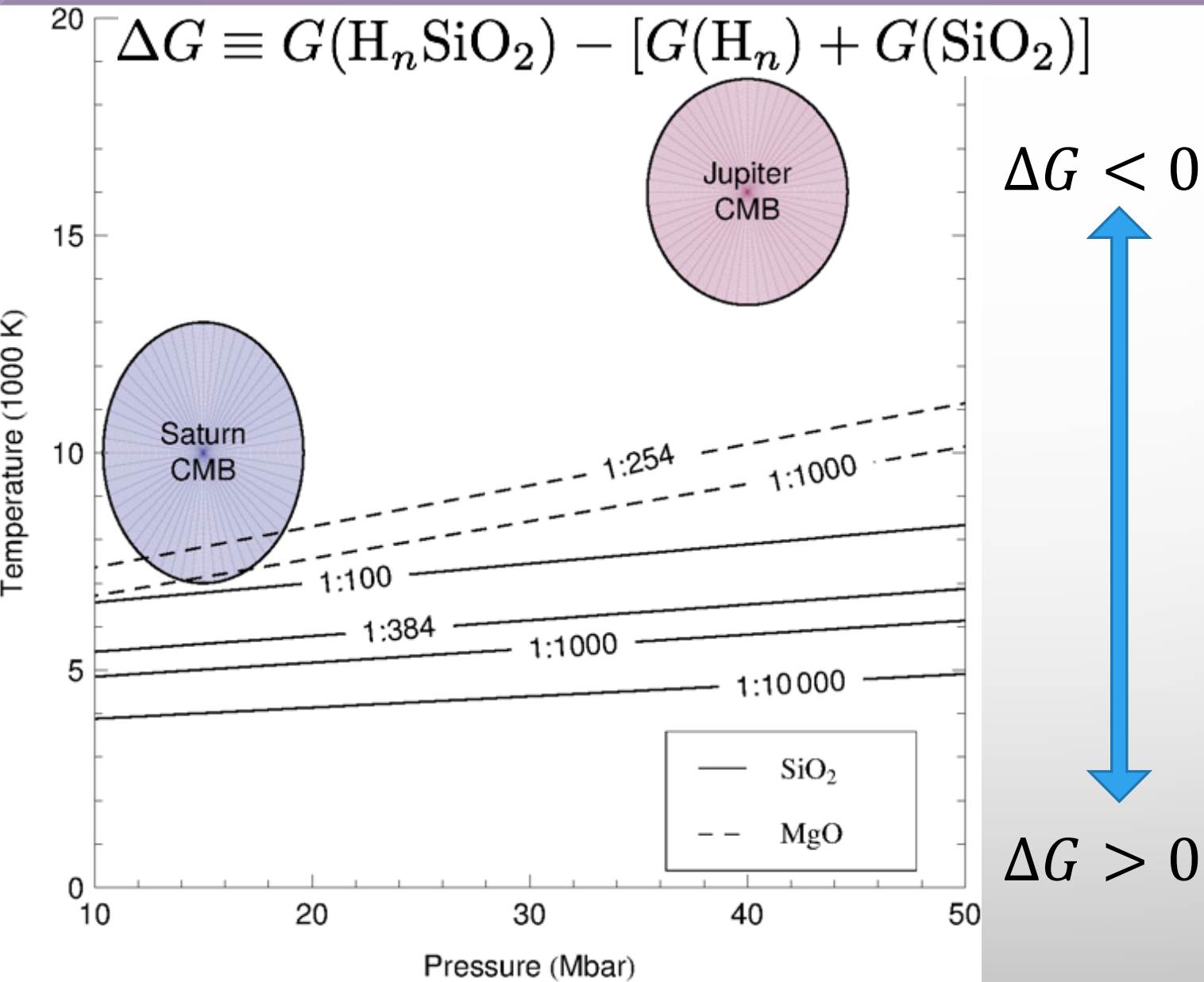


$$U = U_{KS}$$



PLANETARY INTERIORS

A diluted core in Jupiter



$\Delta G < 0$ at CMB

- SiO₂ gets dissolved
- At SiO₂:H < 1:100

- SiO₂ more soluble than MgO
- Fe, H₂O also soluble in H

Wahl+, APJ (2013)

Wilson+ & Militzer APJ (2012)

Gonzalez & Militzer, APJ (2014)

Jupiter's Interior

$$J_n = -\frac{2\pi}{Ma^n} \int dr d\mu \rho(\mathbf{r}) r^{n+2} P_n(\mu)$$

Gravitational
Moments

THE PLANETARY SCIENCE JOURNAL, 3:185 (14pp), 2022 August














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OPEN ACCESS

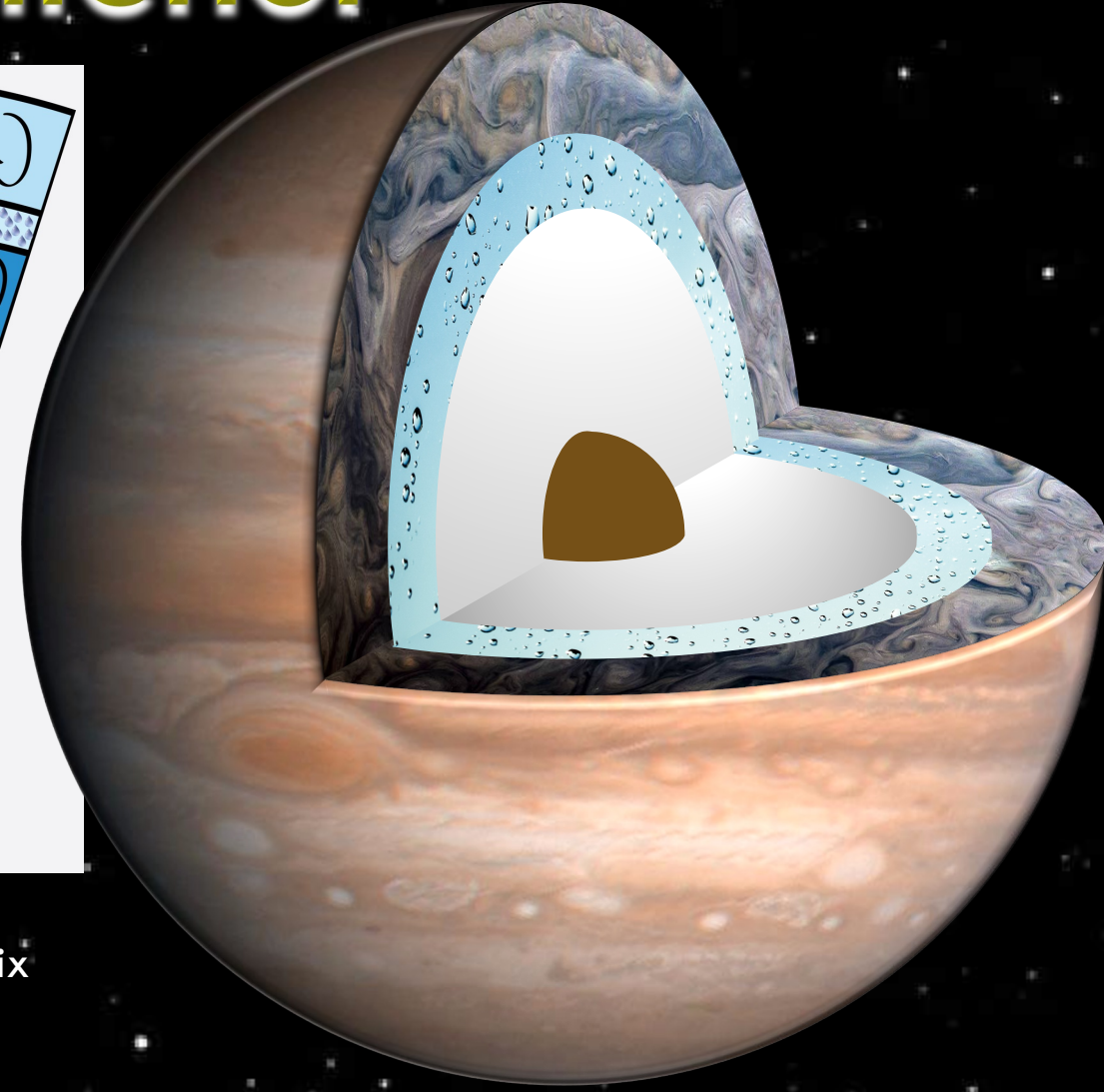
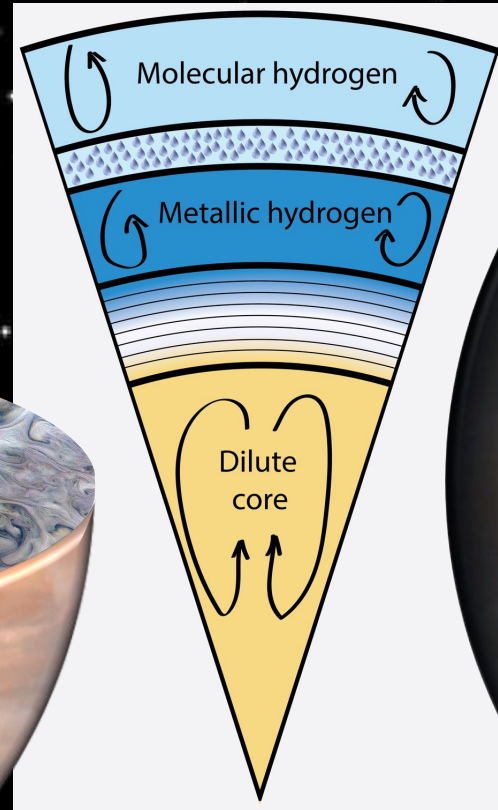
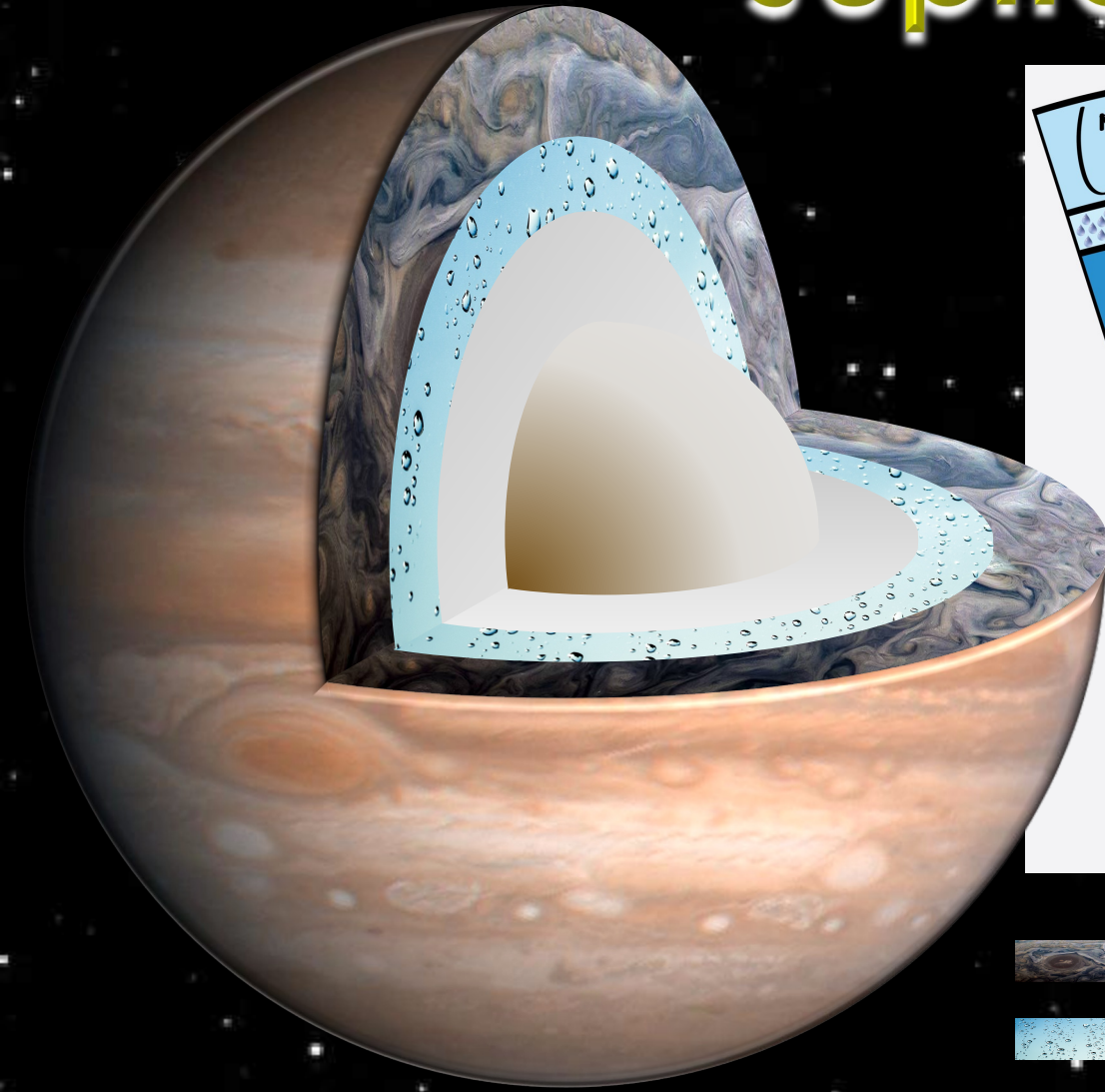
<https://doi.org/10.3847/PSJ/ac7ec8>








Juno Spacecraft Measurements of Jupiter's Gravity Imply a Dilute Core

Burkhard Militzer^{1,2} , William B. Hubbard³ , Sean Wahl¹ , Jonathan I. Lunine⁴ , Eli Galanti⁵ , Yohai Kaspi⁵, Yamila Miguel^{6,7}, Tristan Guillot⁸ , Kimberly M. Moore⁹ , Marzia Parisi¹⁰ , John E. P. Connerney^{11,12}, Ravid Helled¹³ , Hao Cao¹⁴ , Christopher Mankovich⁹ , David J. Stevenson⁹ , Ryan S. Park¹⁰ , Mike Wong^{15,16}, Sushil K. Atreya¹⁷, John Anderson¹⁰, and Scott J. Bolton¹⁸

Jupiter's Interior



-  H + He mix
-  He rain
-  Metallic H
-  Diluted core

 Compact core
(Solid/Liquid?)

OUTLINE

1. Planetary Interiors

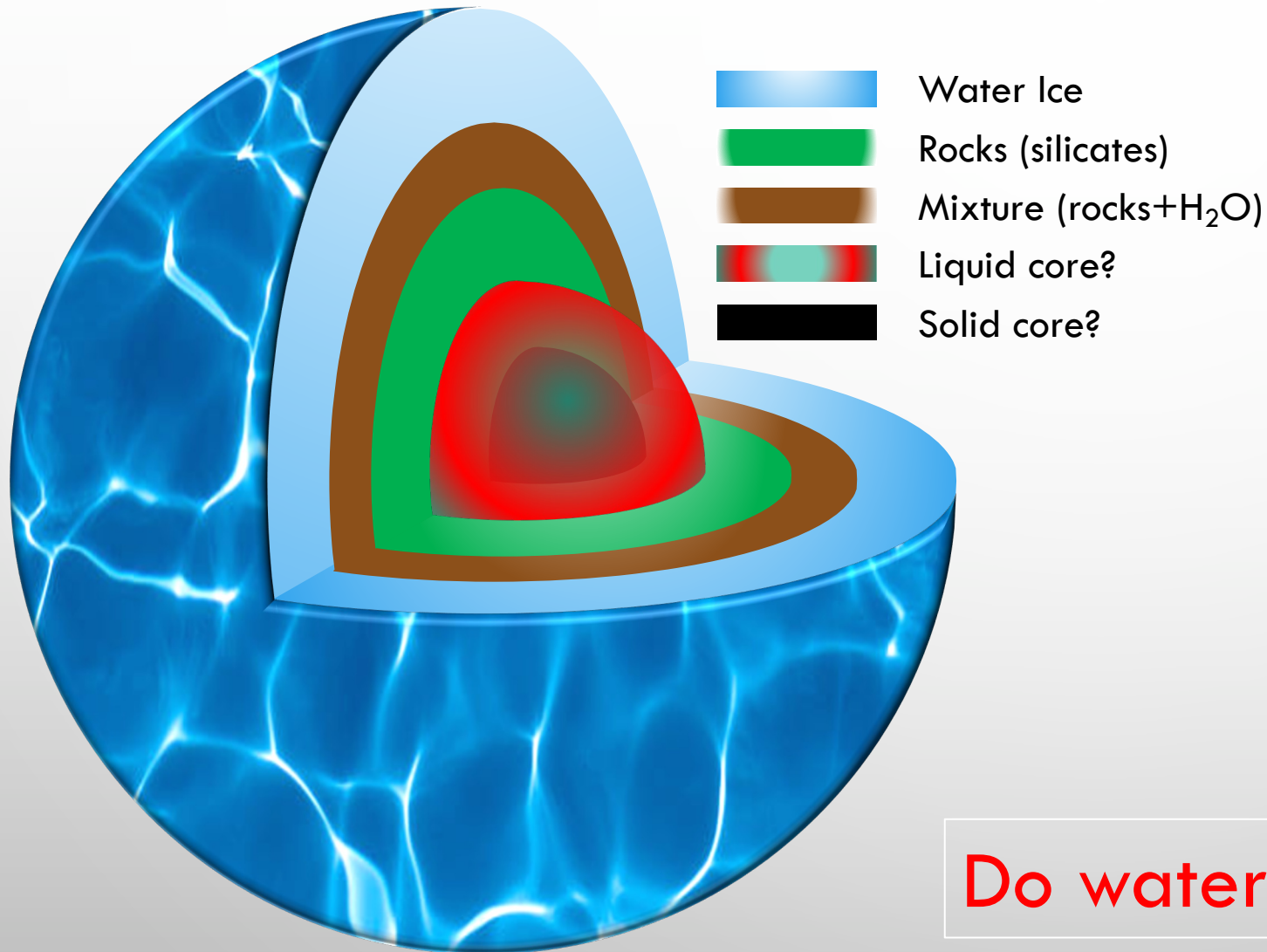
- A diluted core in Jupiter
- Rock/Ice mixtures in water planets

2. High Pressure Phase Transitions

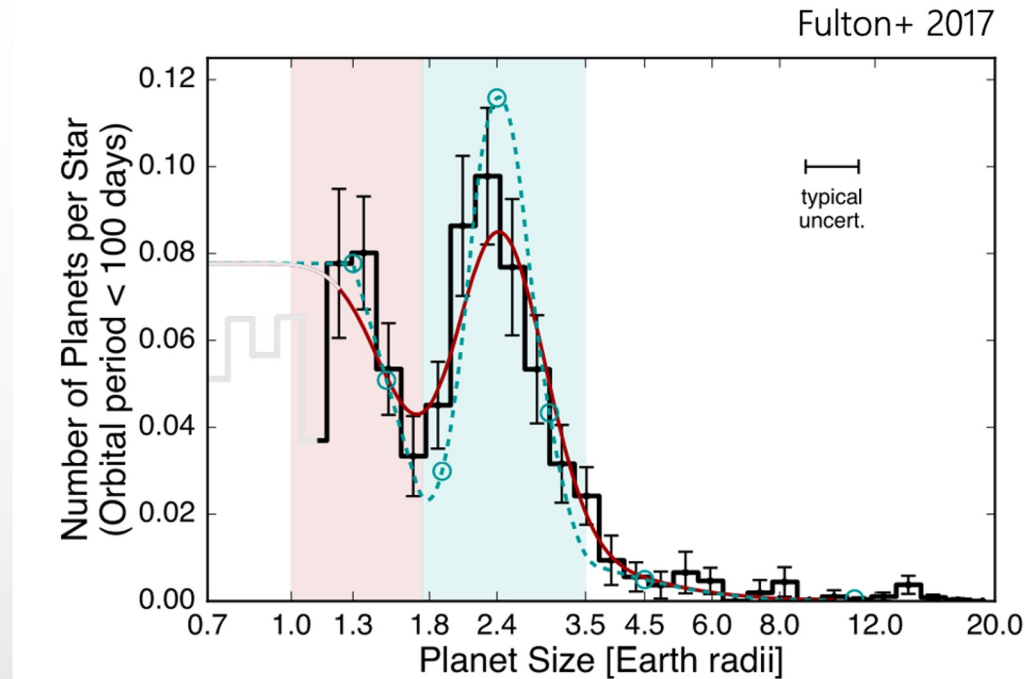
- Be & MgO: melting and anharmonicities
- Melting of SiO_2
- Ramp compression from DFT

3. Warm Dense Matter

- Warm dense silicates: Mg, MgO & MgSiO_3
- FPEOS



Radius Gap (Fulton Gap)

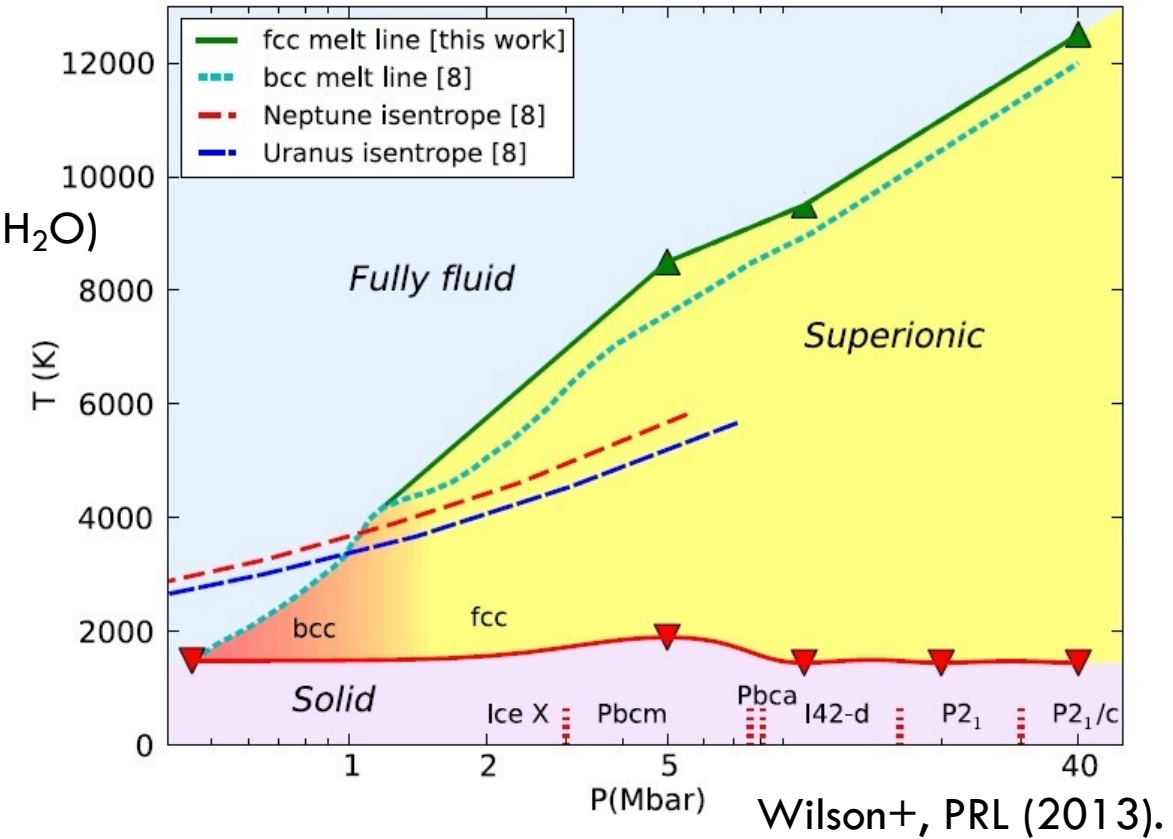
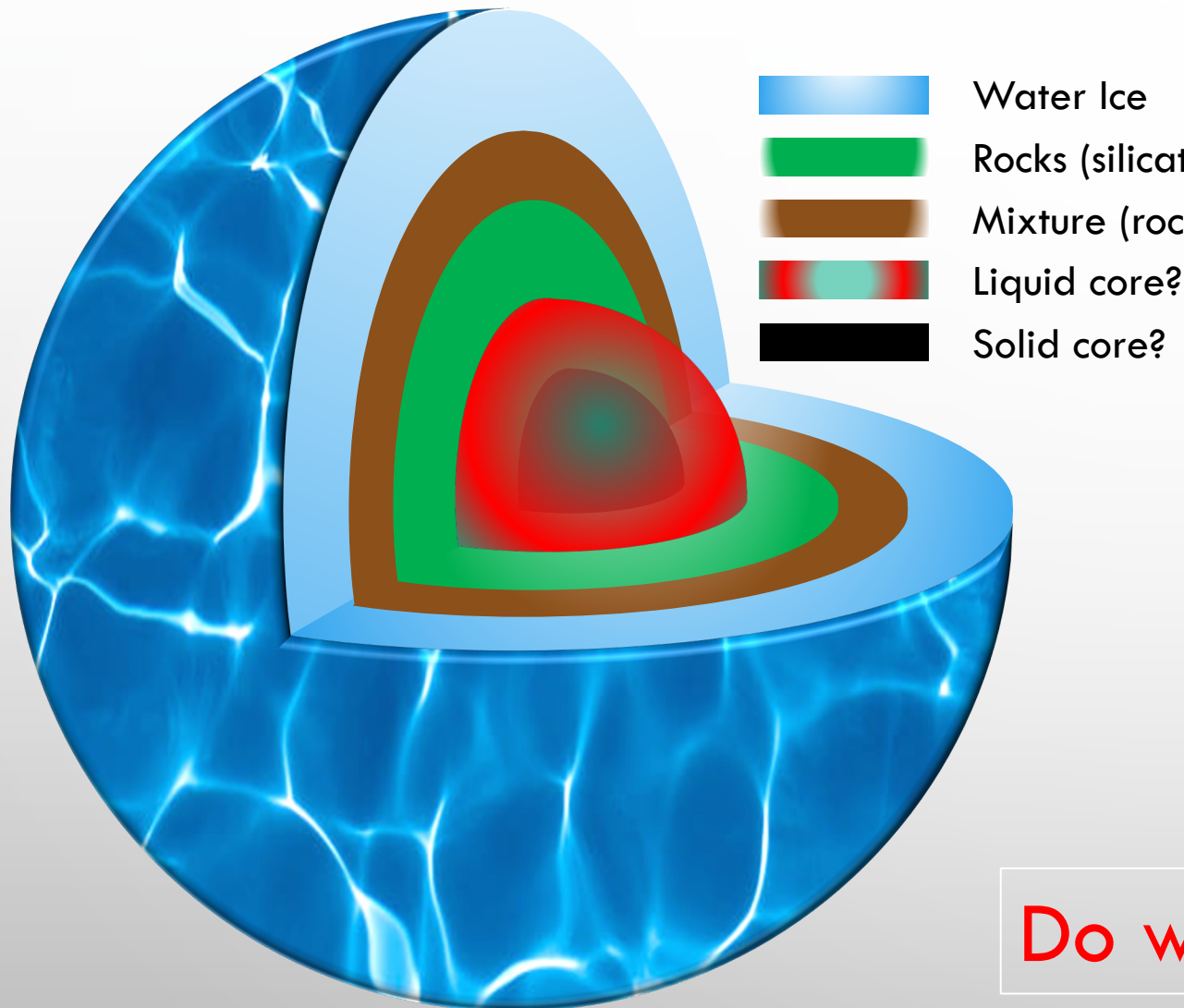


Super-Earths → Sub-Neptunes

Do water and silicates mix?

PLANETARY INTERIORS

WATER WORLDS



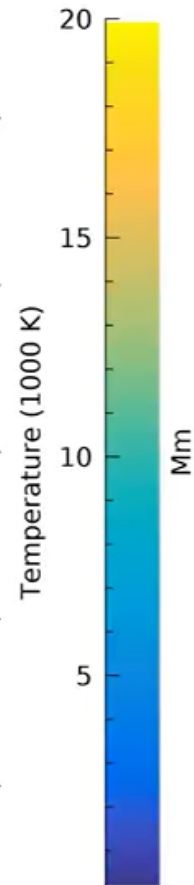
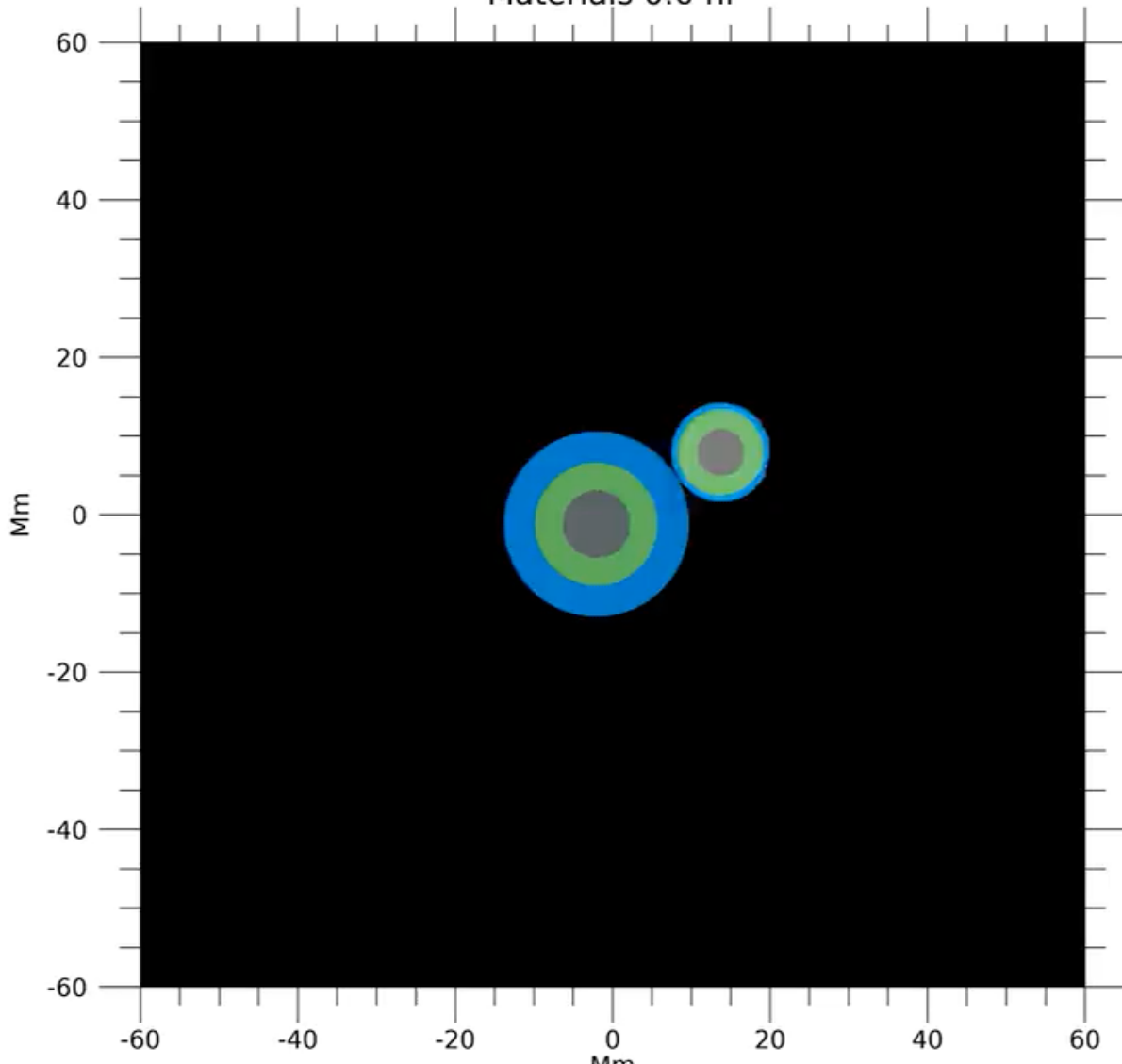
Do water and silicates mix?

PLANETARY INTERIORS

WATER WORLDS

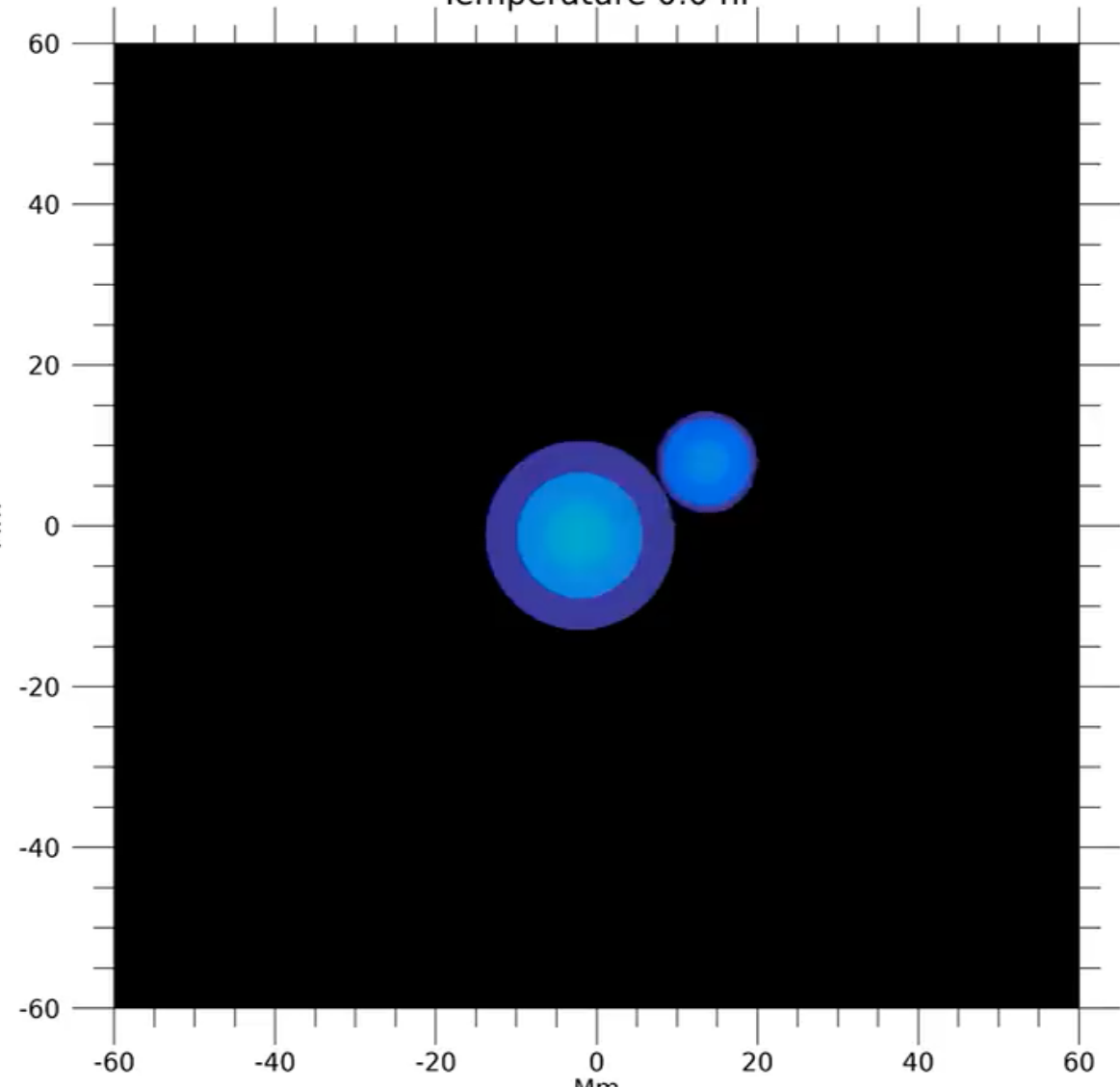
Smooth Particle Hydrodynamics

Materials 0.0 hr



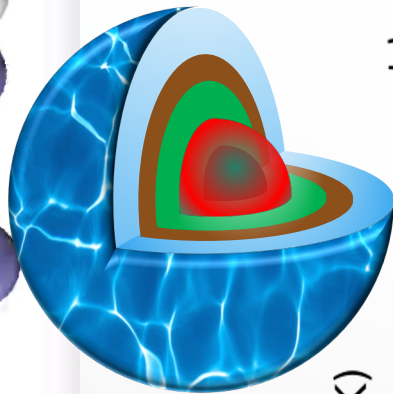
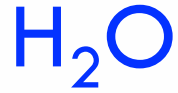
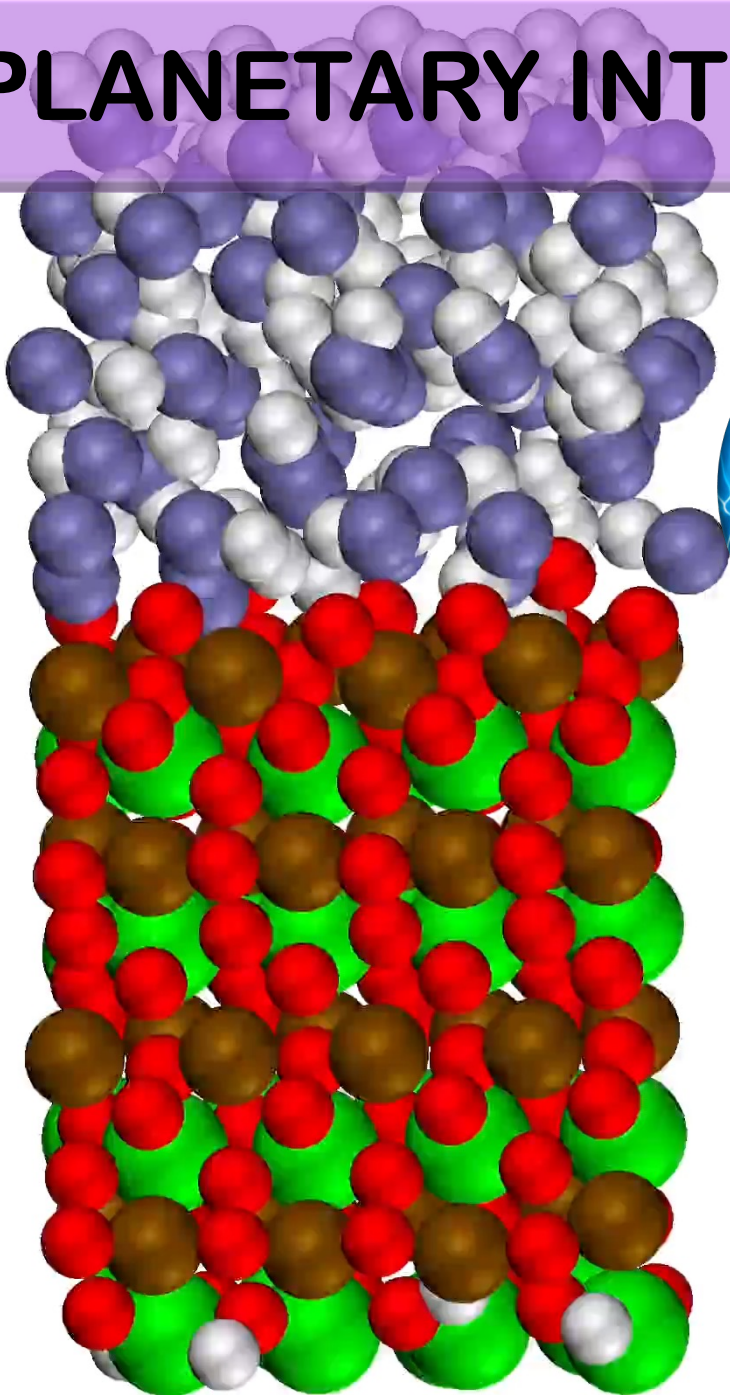
Kovacevic+, Sci. Rep. (2022)

Temperature 0.0 hr



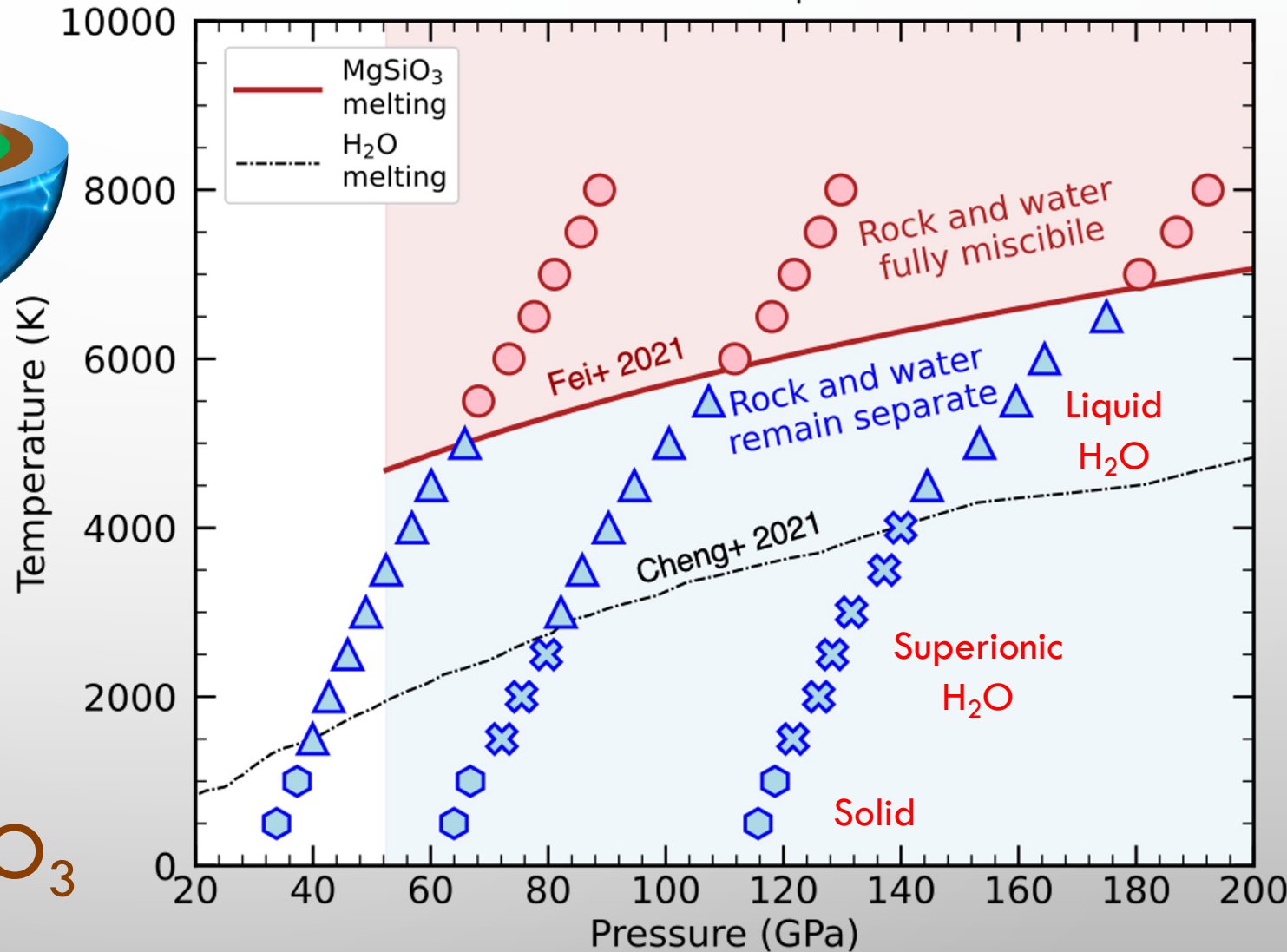
PLANETARY INTERIORS

WATER WORLDS



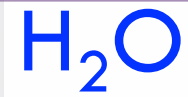
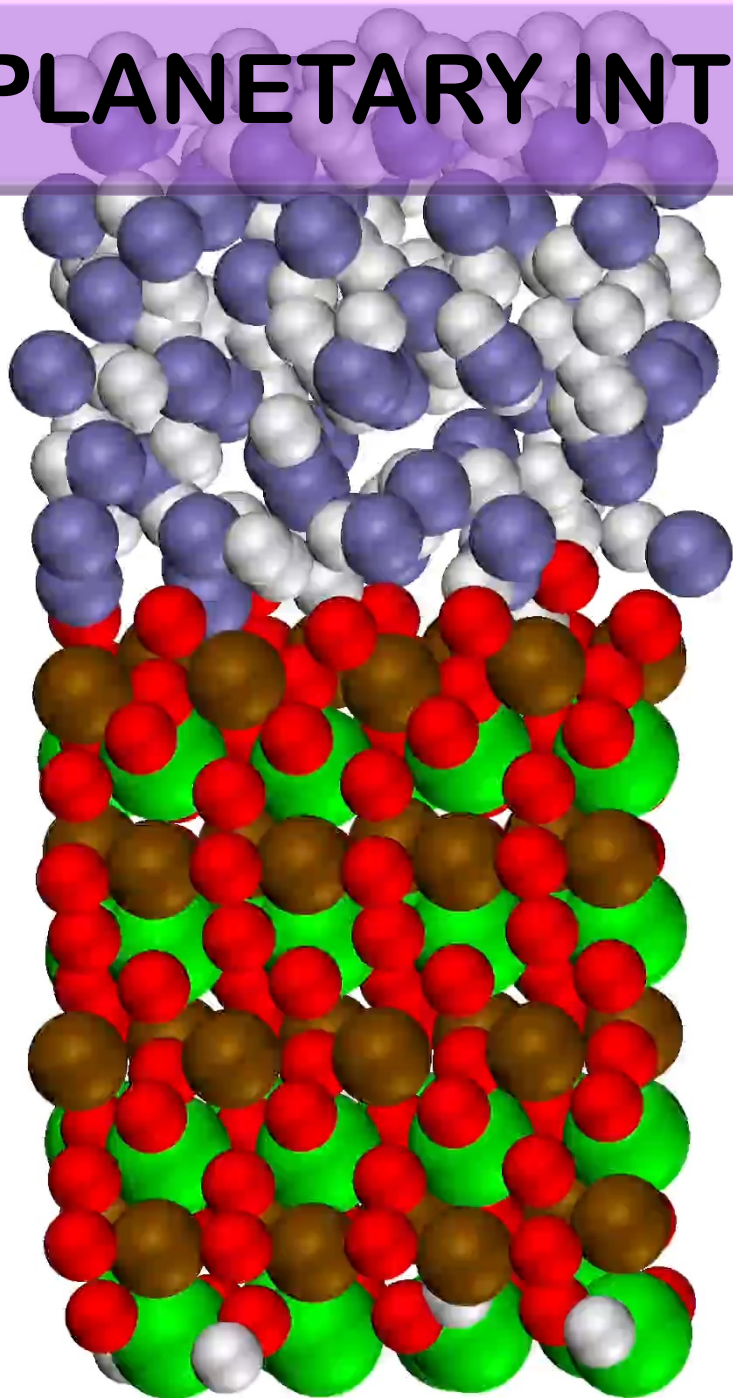
Kovacevic+, Sci. Rep. (2022)

Adapted from Kovacevic+ 2022

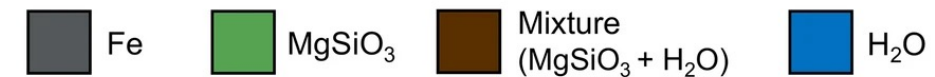
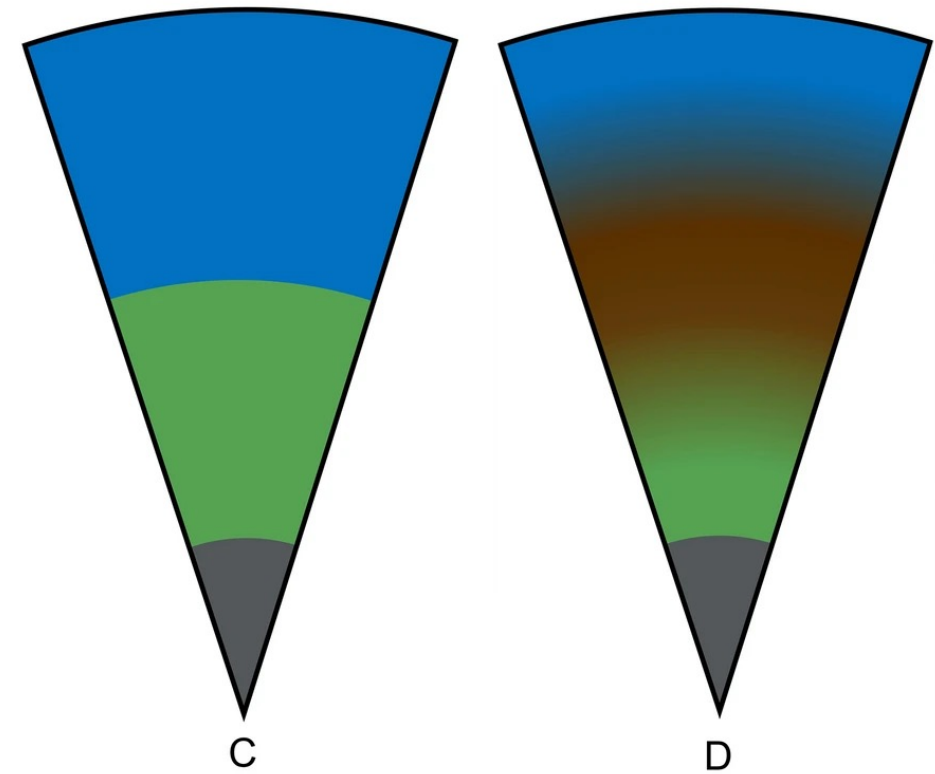
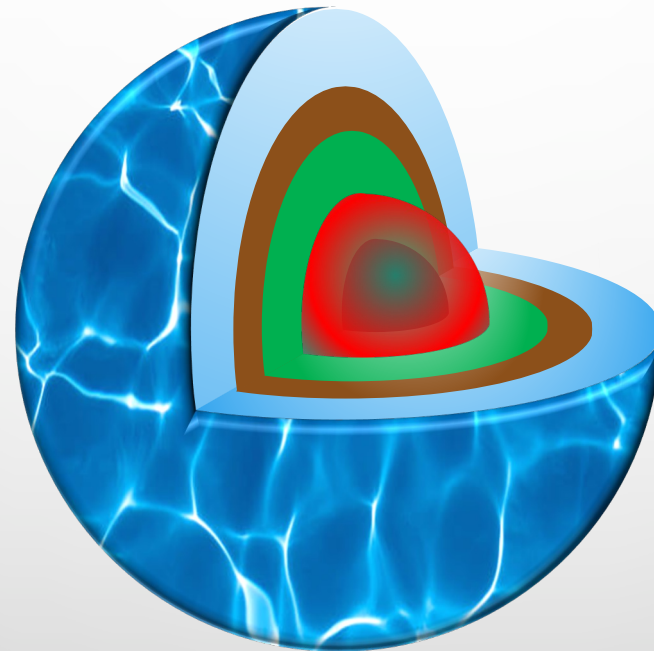


PLANETARY INTERIORS

WATER WORLDS



Kovacevic+, Sci. Rep. (2022)



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- A diluted core in Jupiter
- Rock/Ice mixtures in water planets

2. High Pressure Phase Transitions

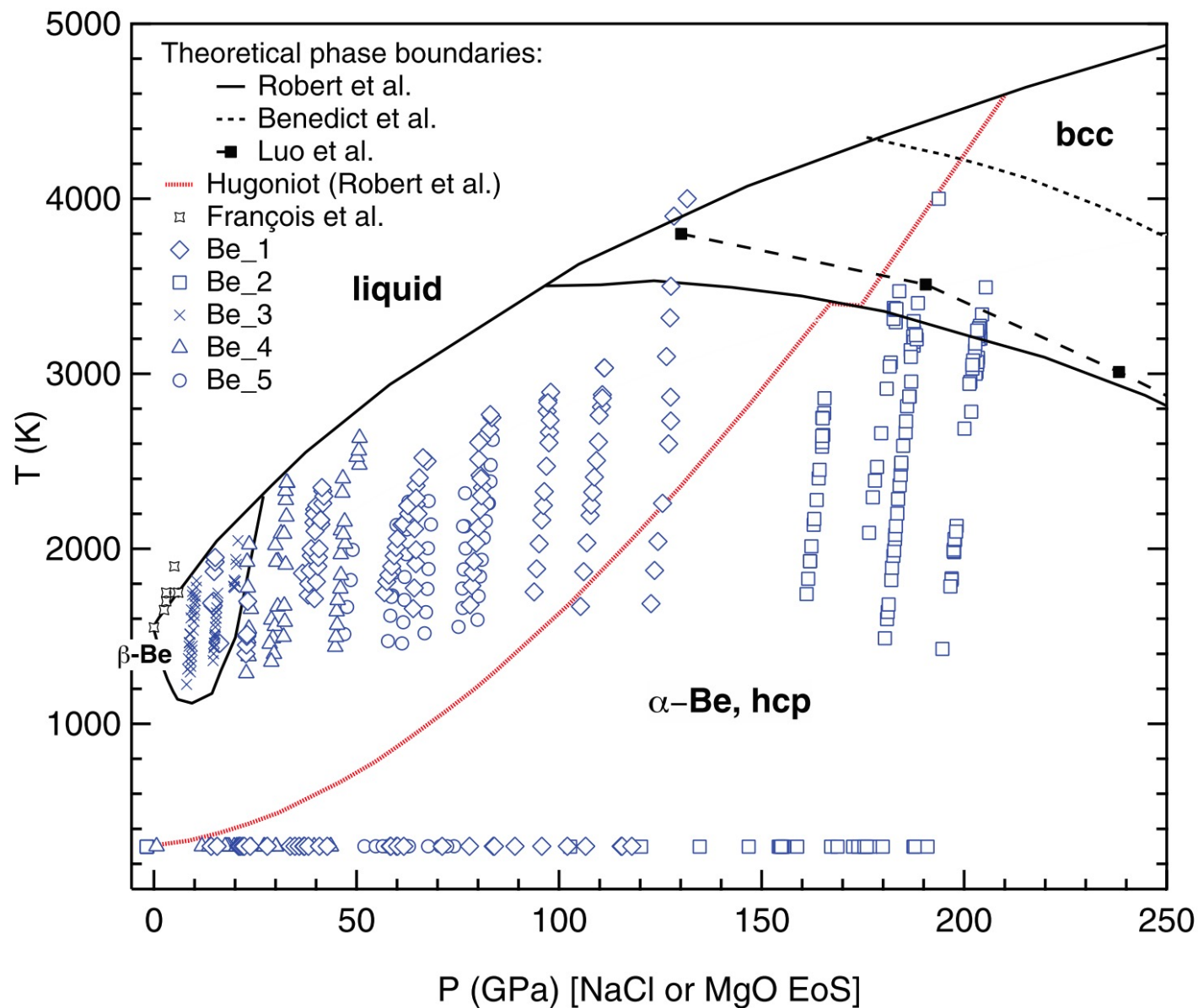
- Be & MgO: melting and anharmonicities
- Melting of SiO_2
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3. Warm Dense Matter

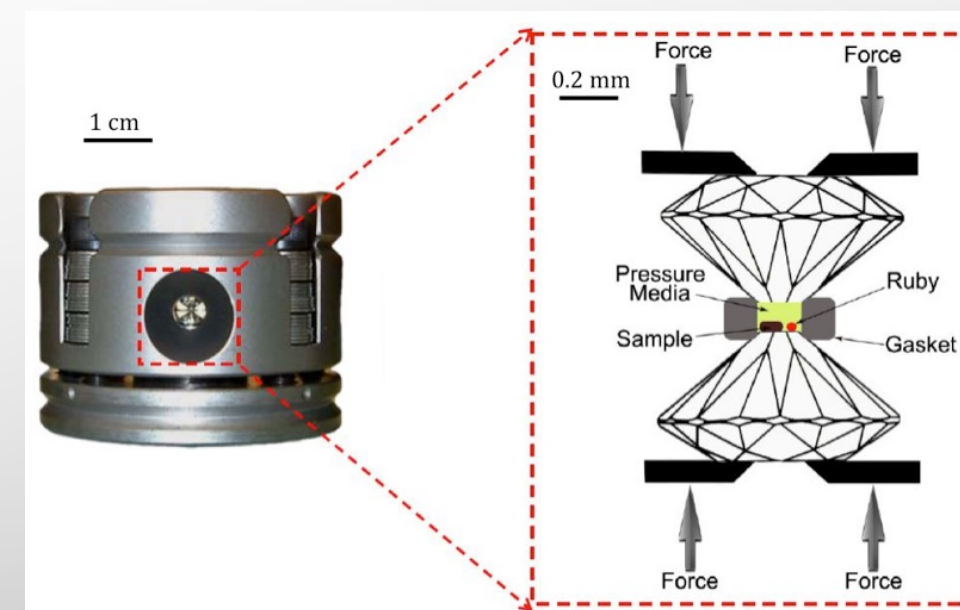
- Warm dense silicates: Mg, MgO & MgSiO_3
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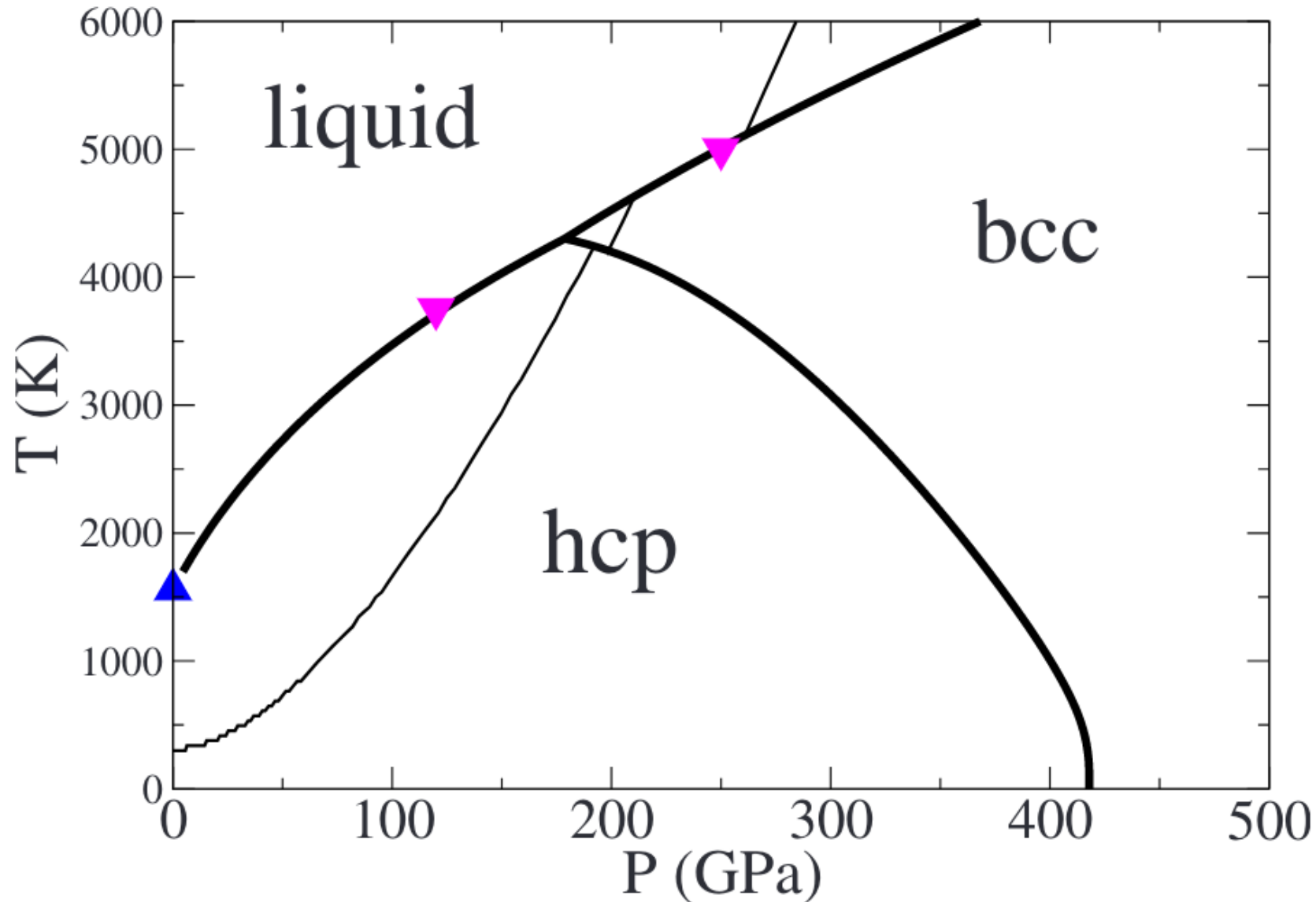
HIGH PRESSURE

Beryllium



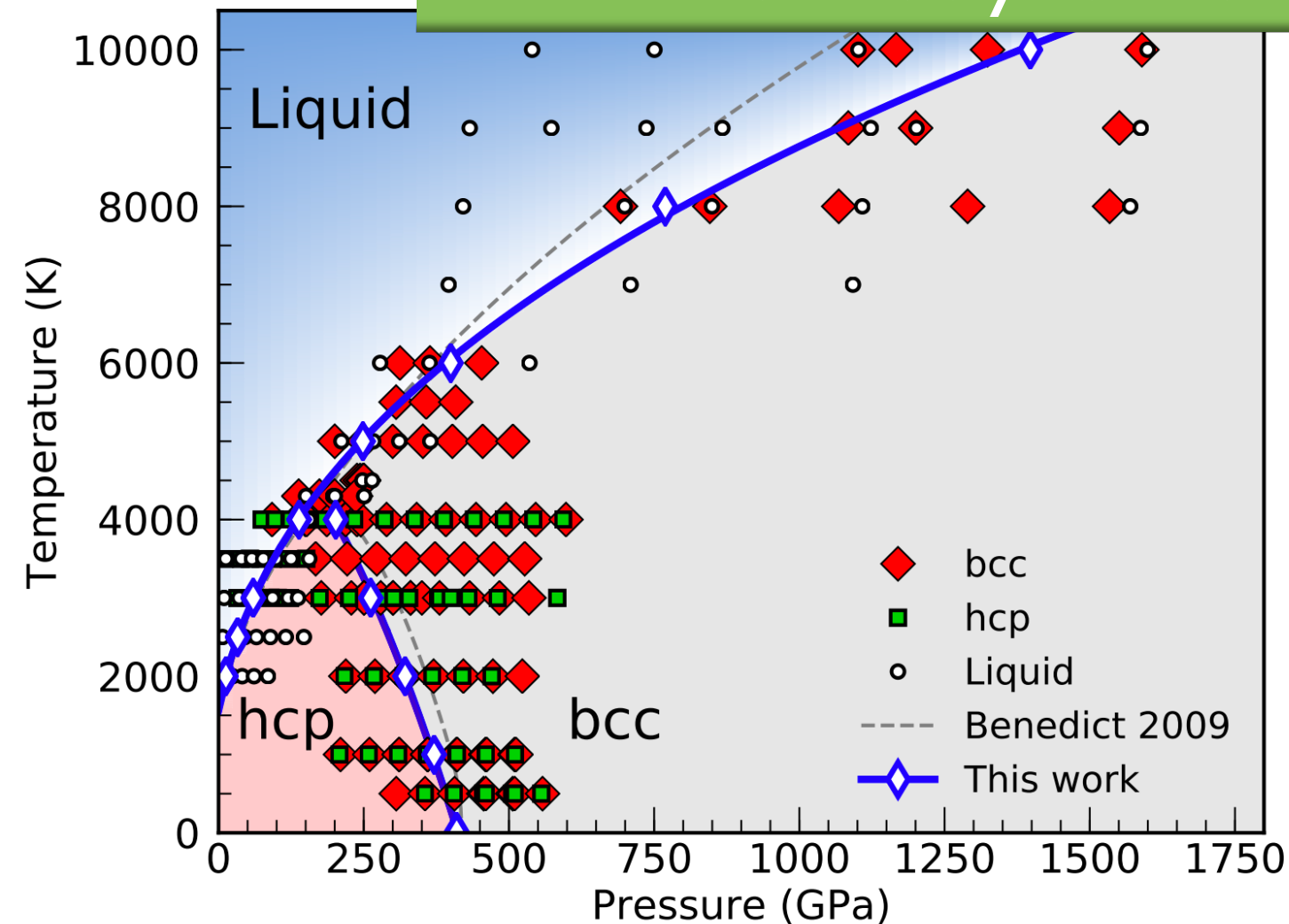
- Laser Heated DAC
- No signature of bcc



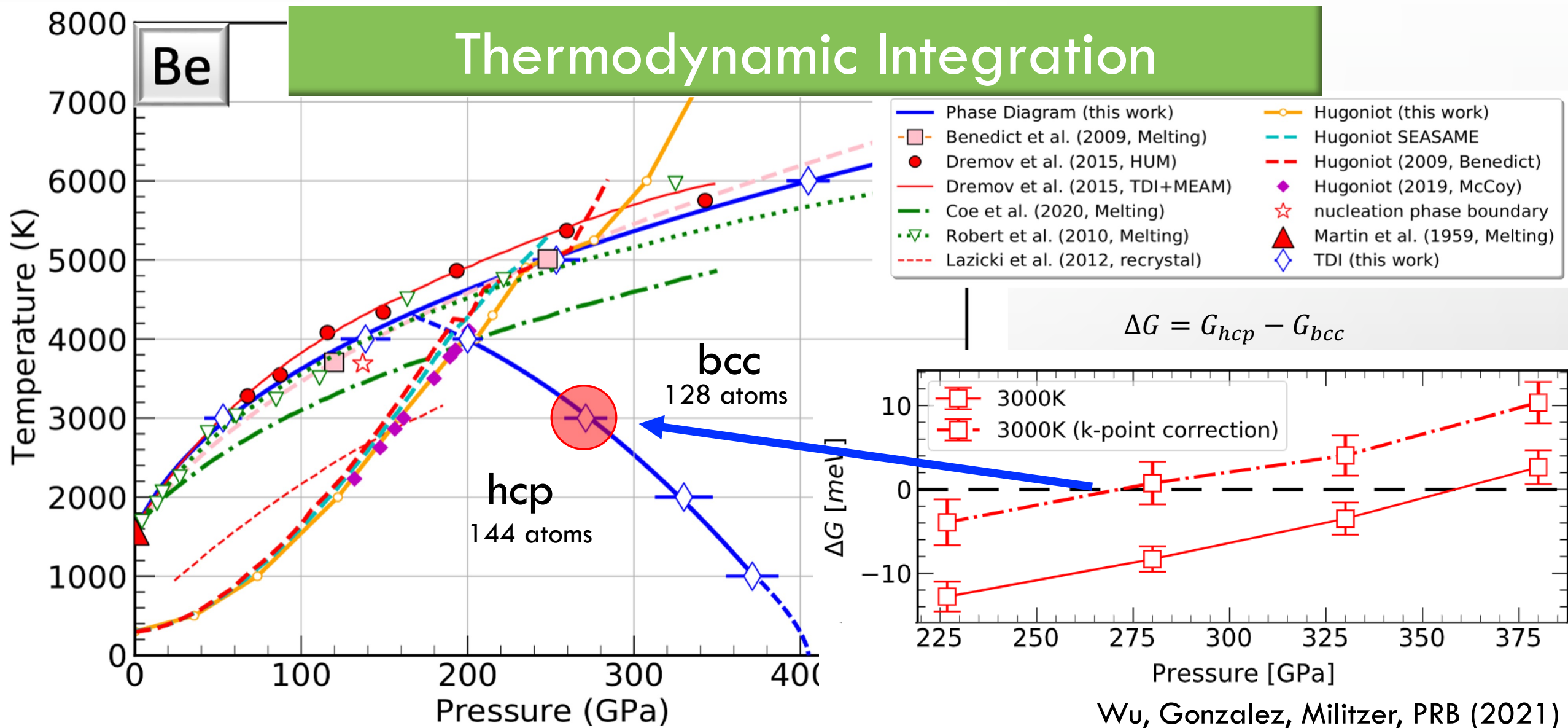


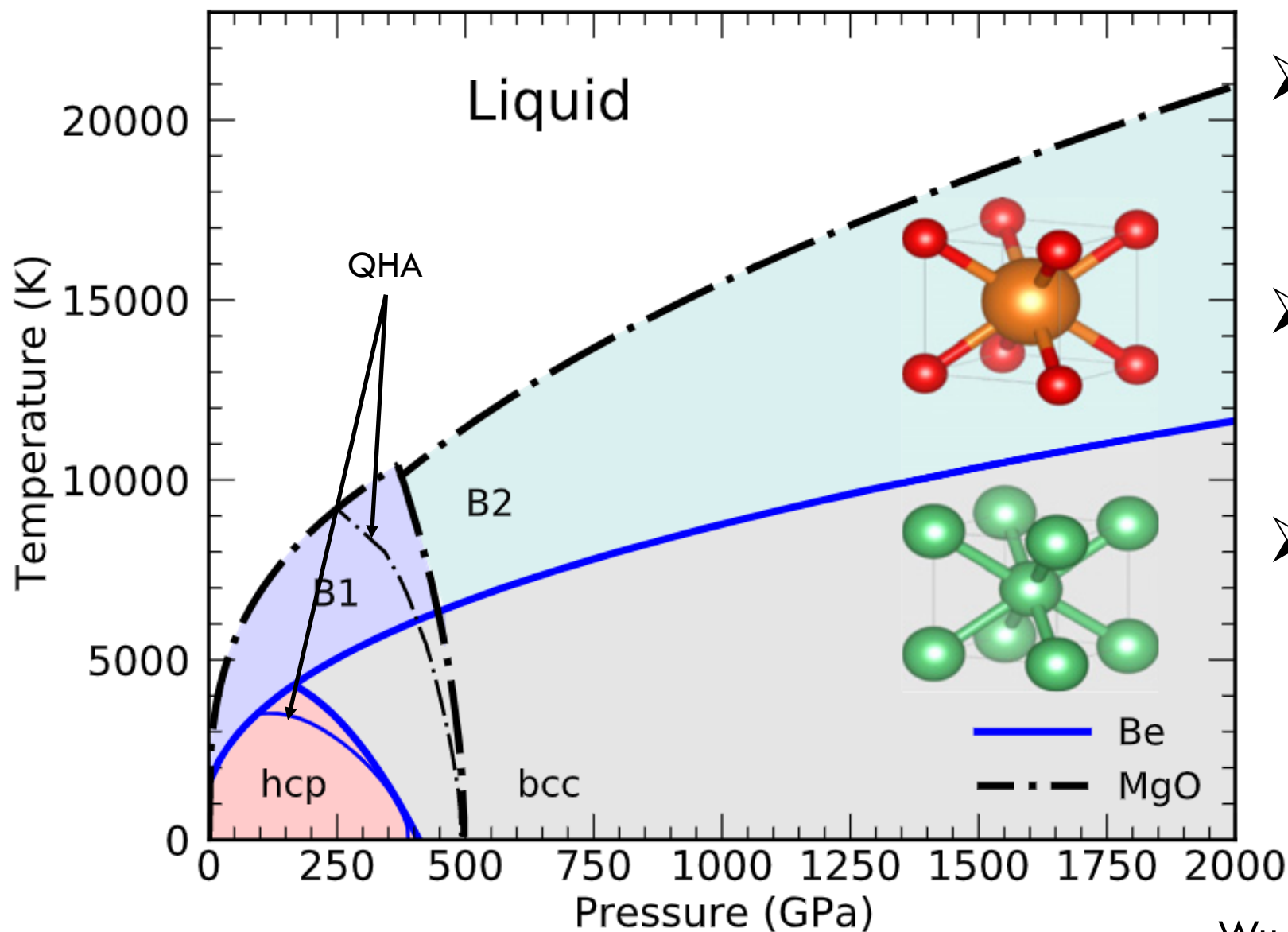
- QHA (quasi-harmonic approximation)
- No signature of bcc in experiments
- Two-phase simulations of melting

Thermodynamic Integration



- DFT-MD
- hcp/bcc/liquid



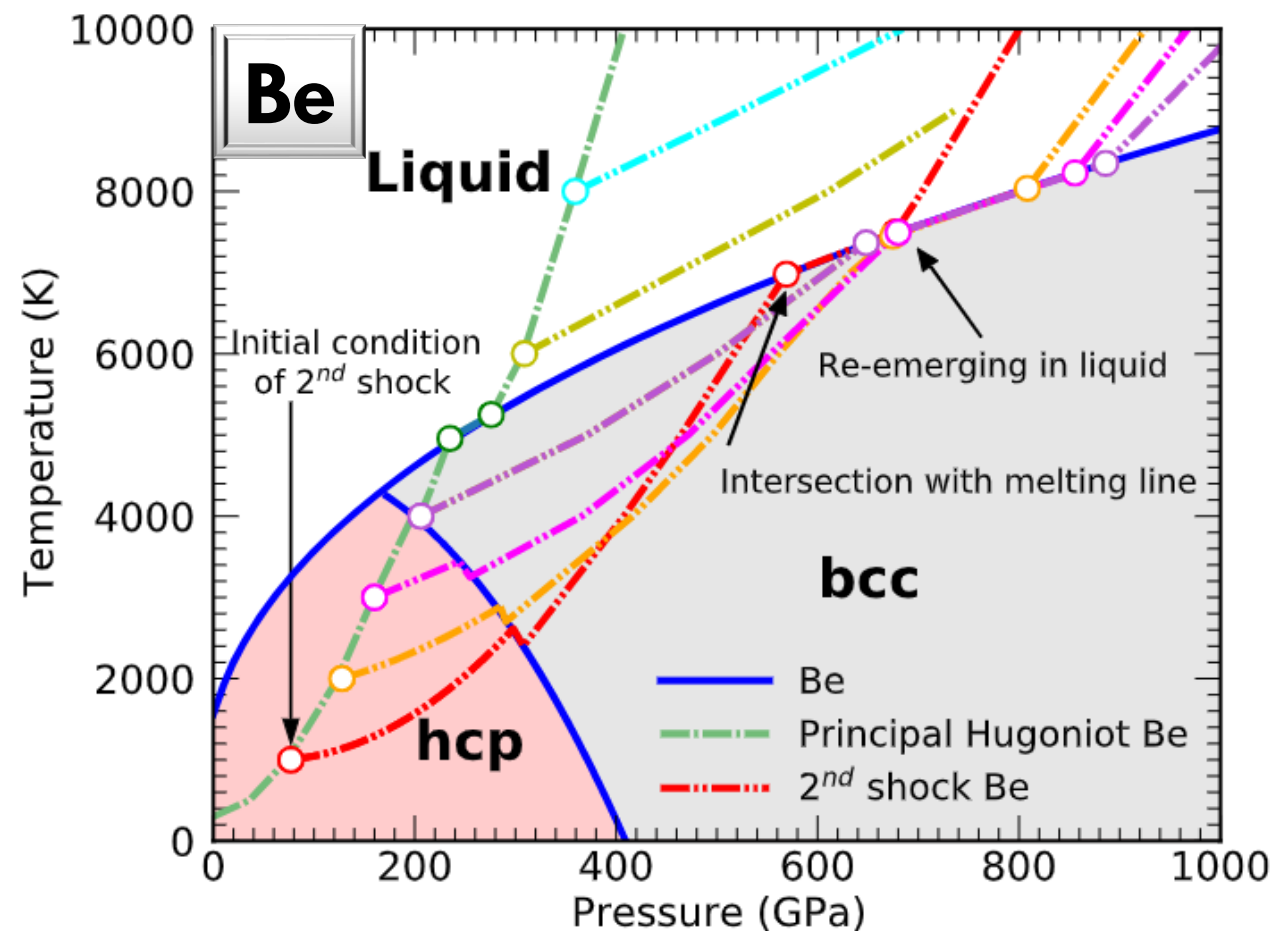
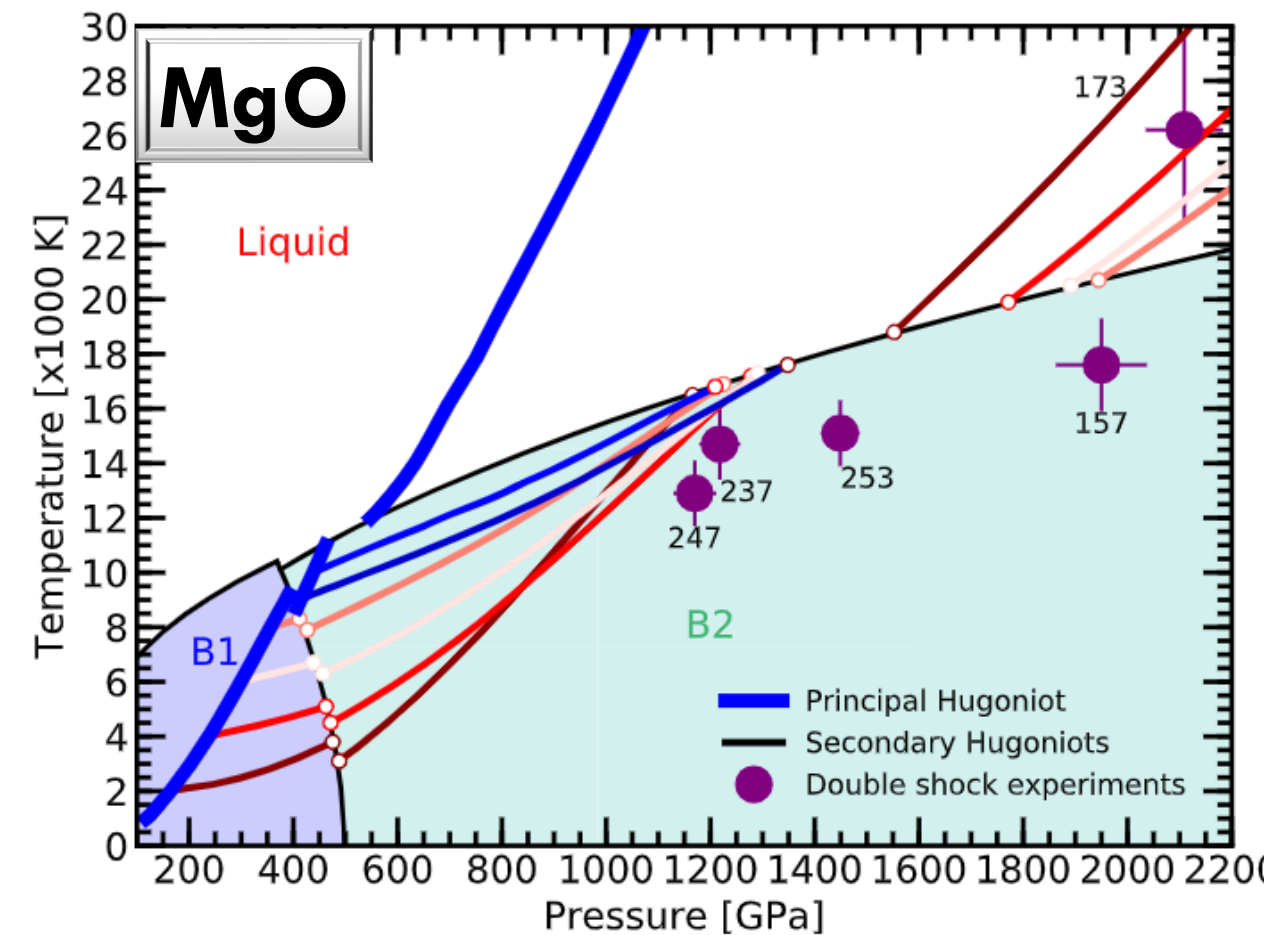


- $\Delta G = G_{liq} - G_{sol}$
Melting lines Be & MgO
- 2x2x2 k-points in Be in
~100-atoms cells
- Strong anharmonicities:
 - QHA does not work well.
 - B1-B2 / hcp-bcc needs higher P.

HIGH PRESSURE

Beryllium & MgO

Double Shock compression



OUTLINE

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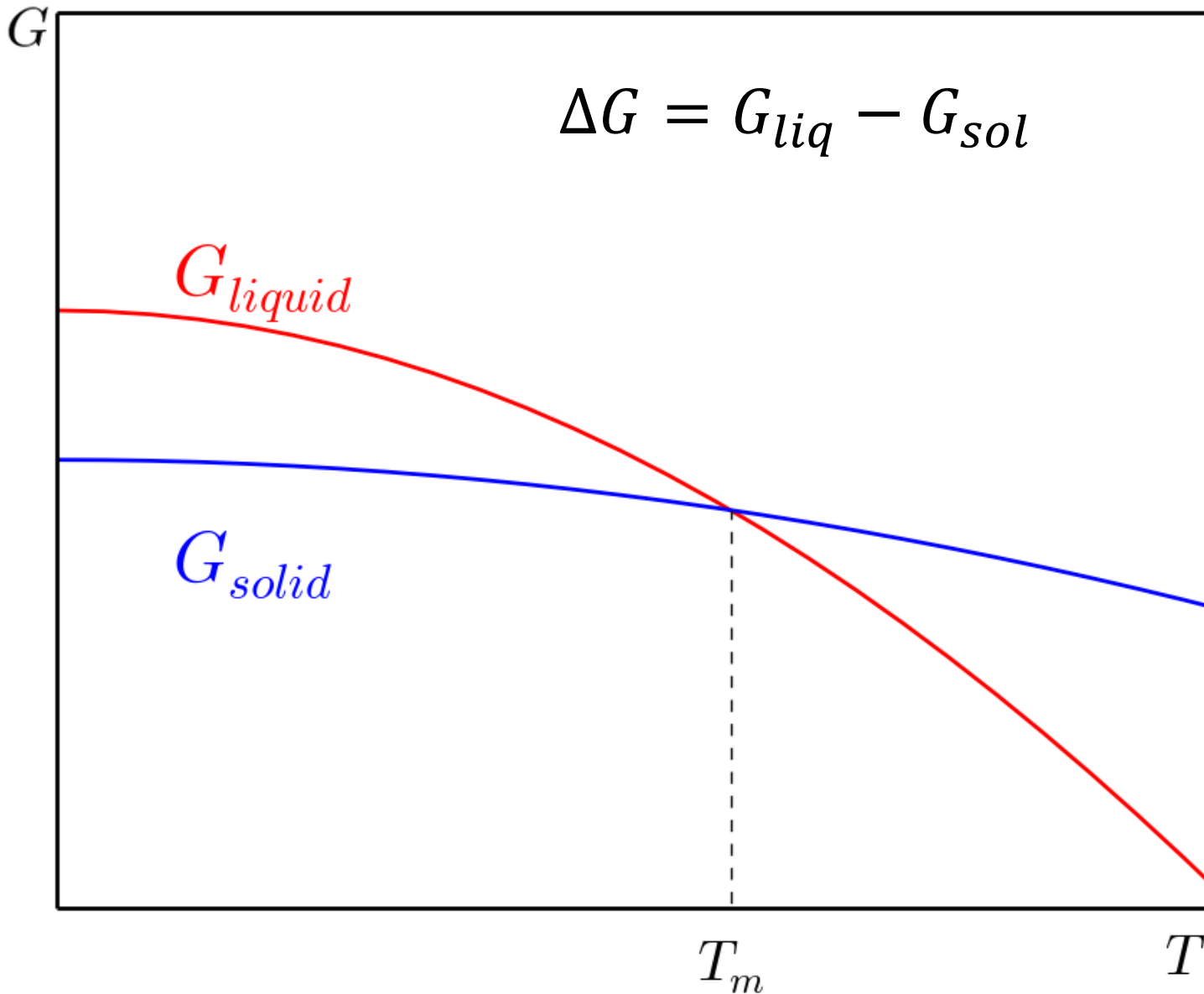
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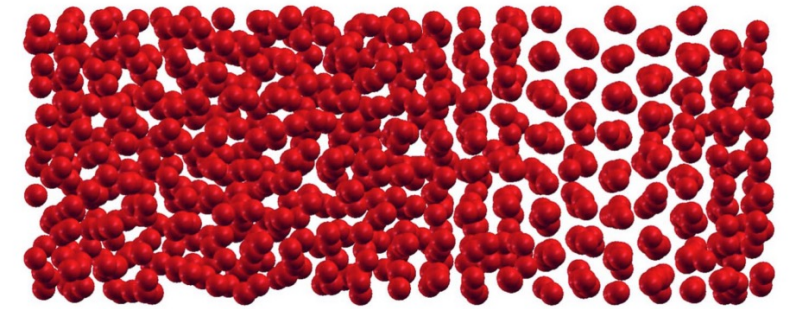
- Warm dense silicates: Mg, MgO & MgSiO_3
- FPEOS

HIGH PRESSURE

Melting SiO_2

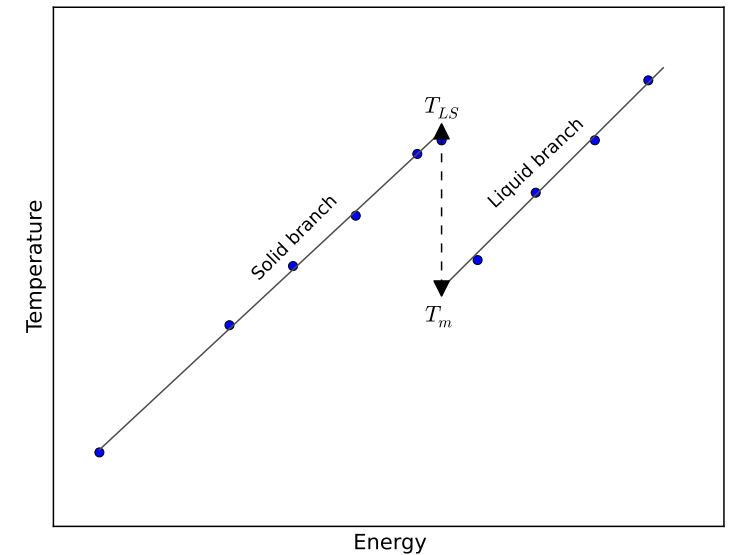


Coexistence (Two phase)



Alfe, PRB (2009)

Z method

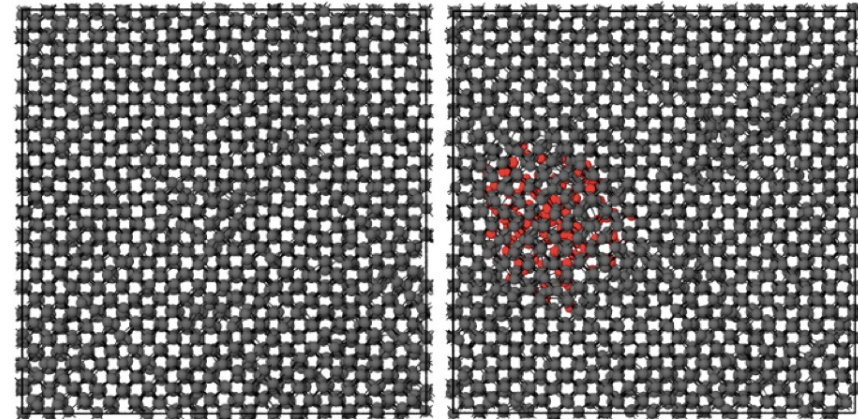
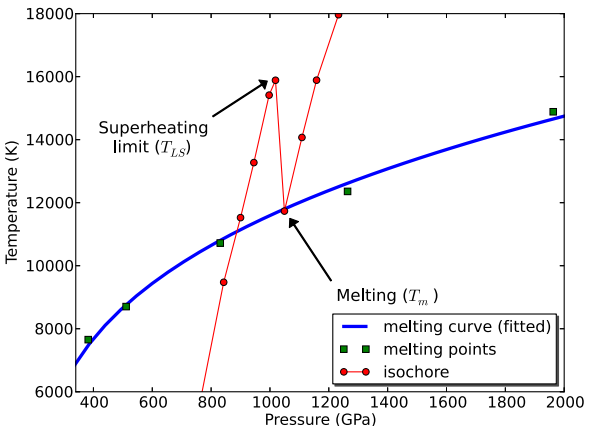
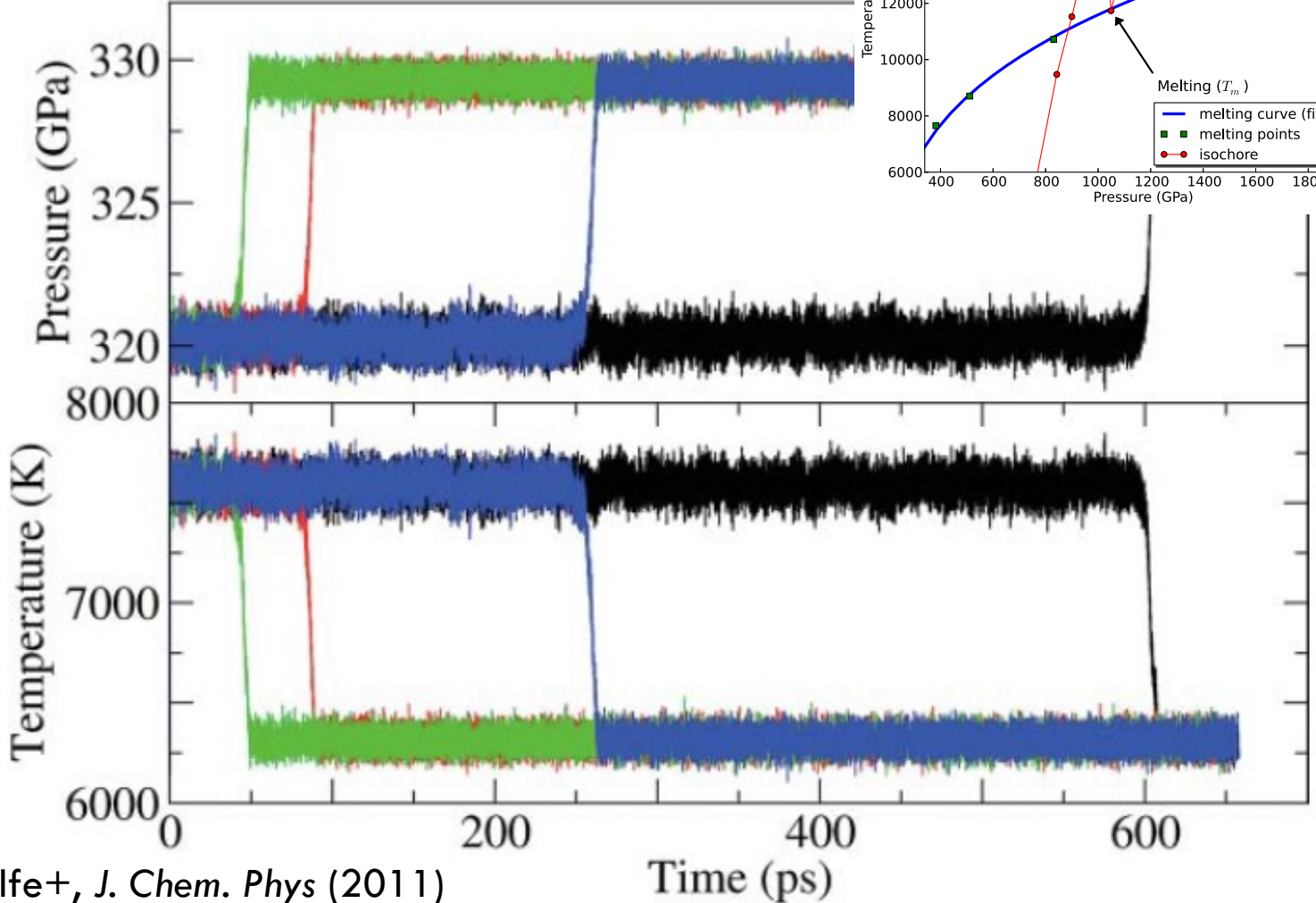


HIGH PRESSURE

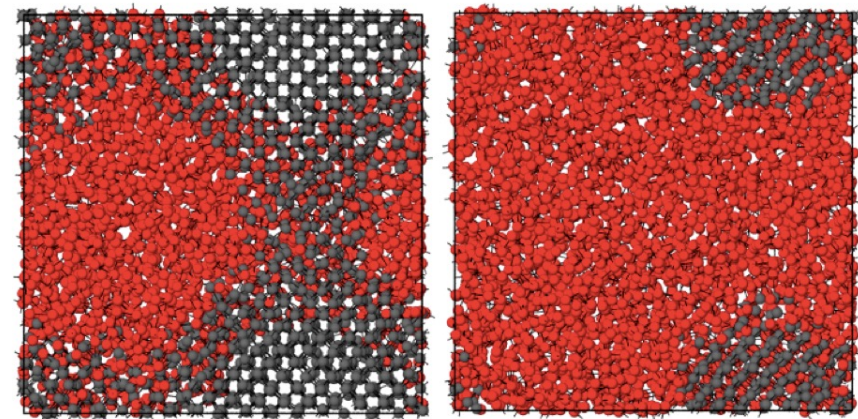
Melting SiO_2

Z method

NVE heat until it melts



(a) $t = 0.50ps$ (b) $t = 0.75ps$



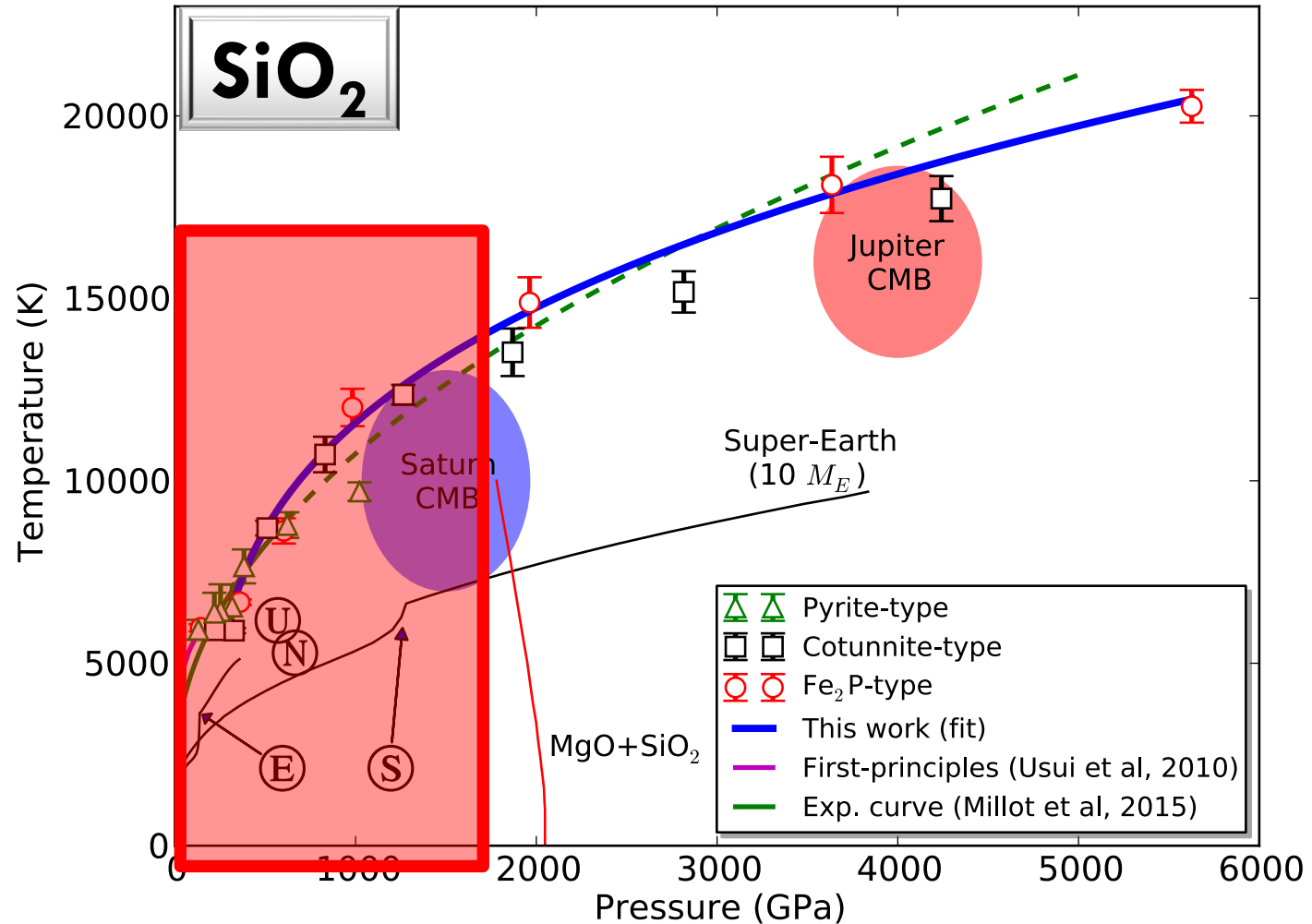
(c) $t = 1.50ps$ (d) $t = 4.00ps$

Gonzalez+*J. Phys. Conf. Series* (2018)

Alfe+, *J. Chem. Phys* (2011)

Z method

NVE heat until it melts

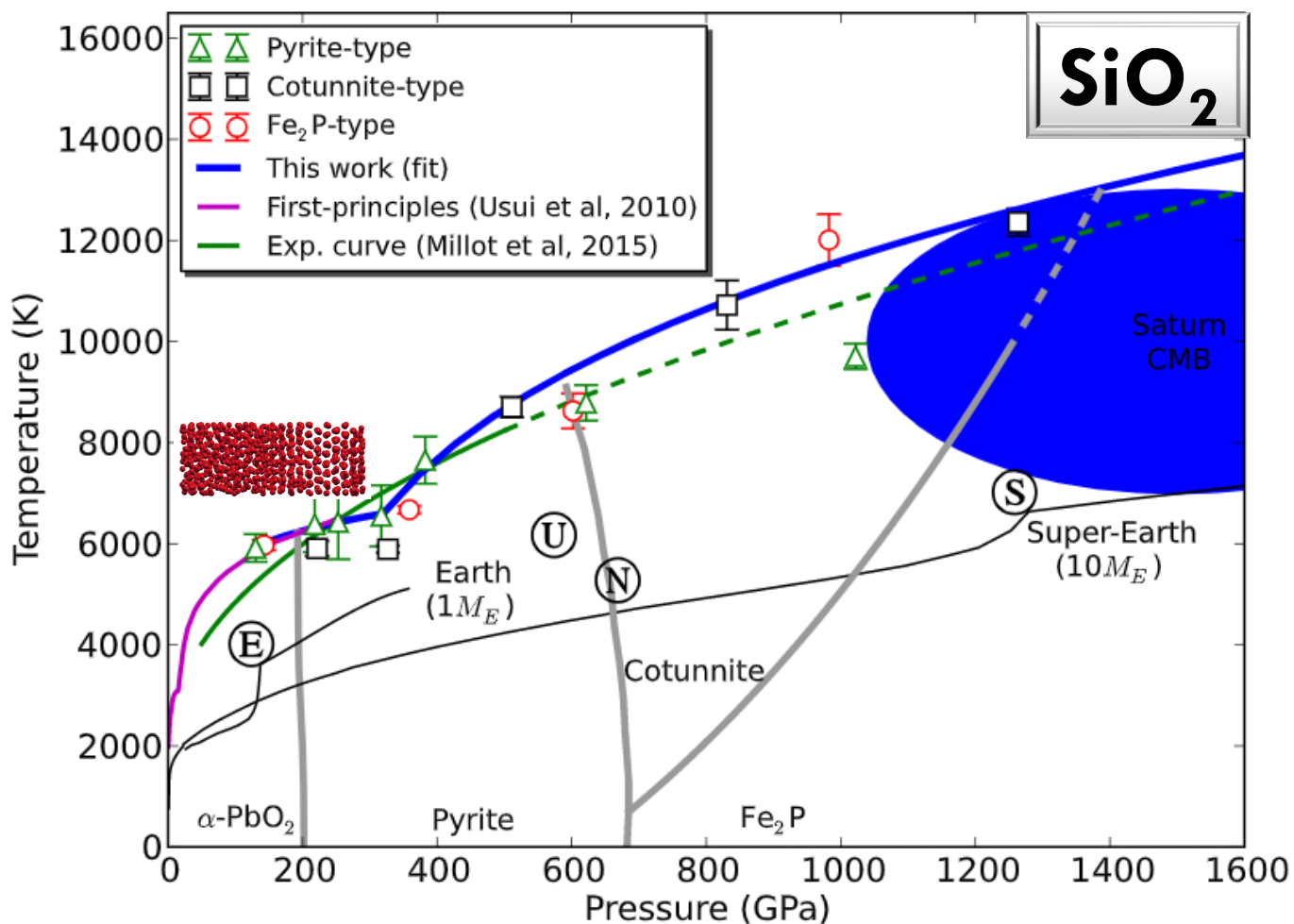


- SiO_2 is solid at CMB of
- Super-Earths
- Gas Giants

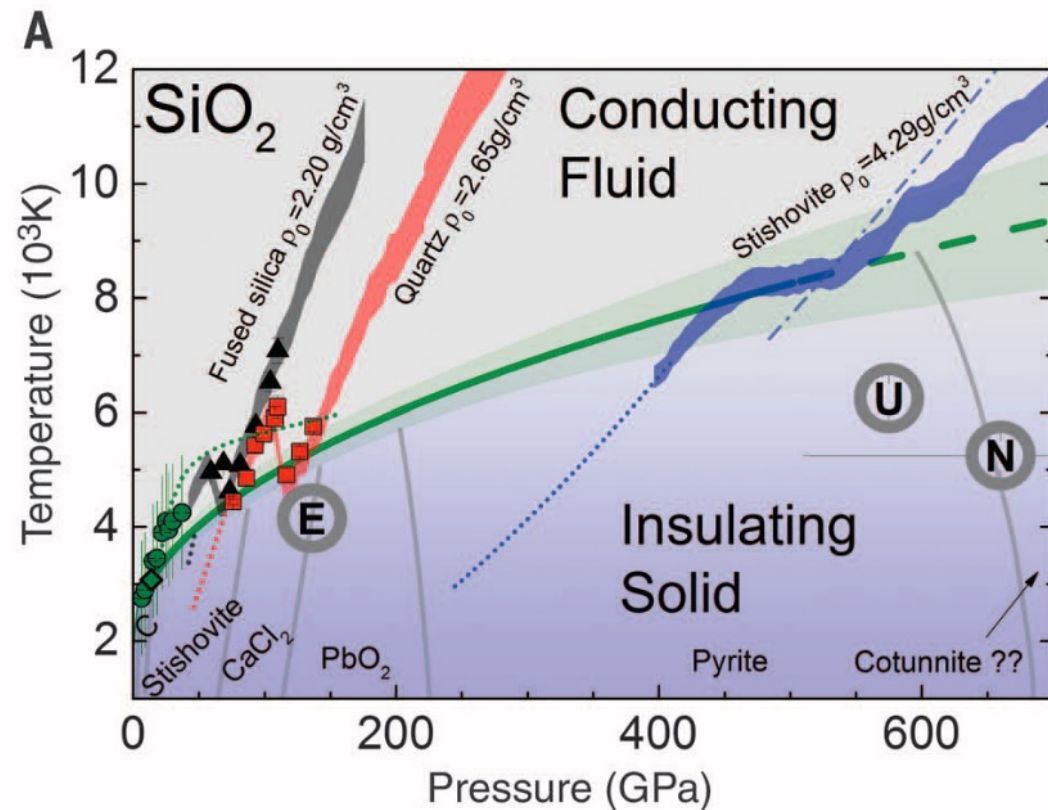
HIGH PRESSURE

Z method

NVE heat until it melts



Melting SiO₂



Millot+, Science (2015)

Z method agreement with

- Two-phase simulations
- Experiments

Gonzalez, Davis, Gutierrez, Sci. Rep. (2015)

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- Ramp compression from DFT

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Ramp compression model

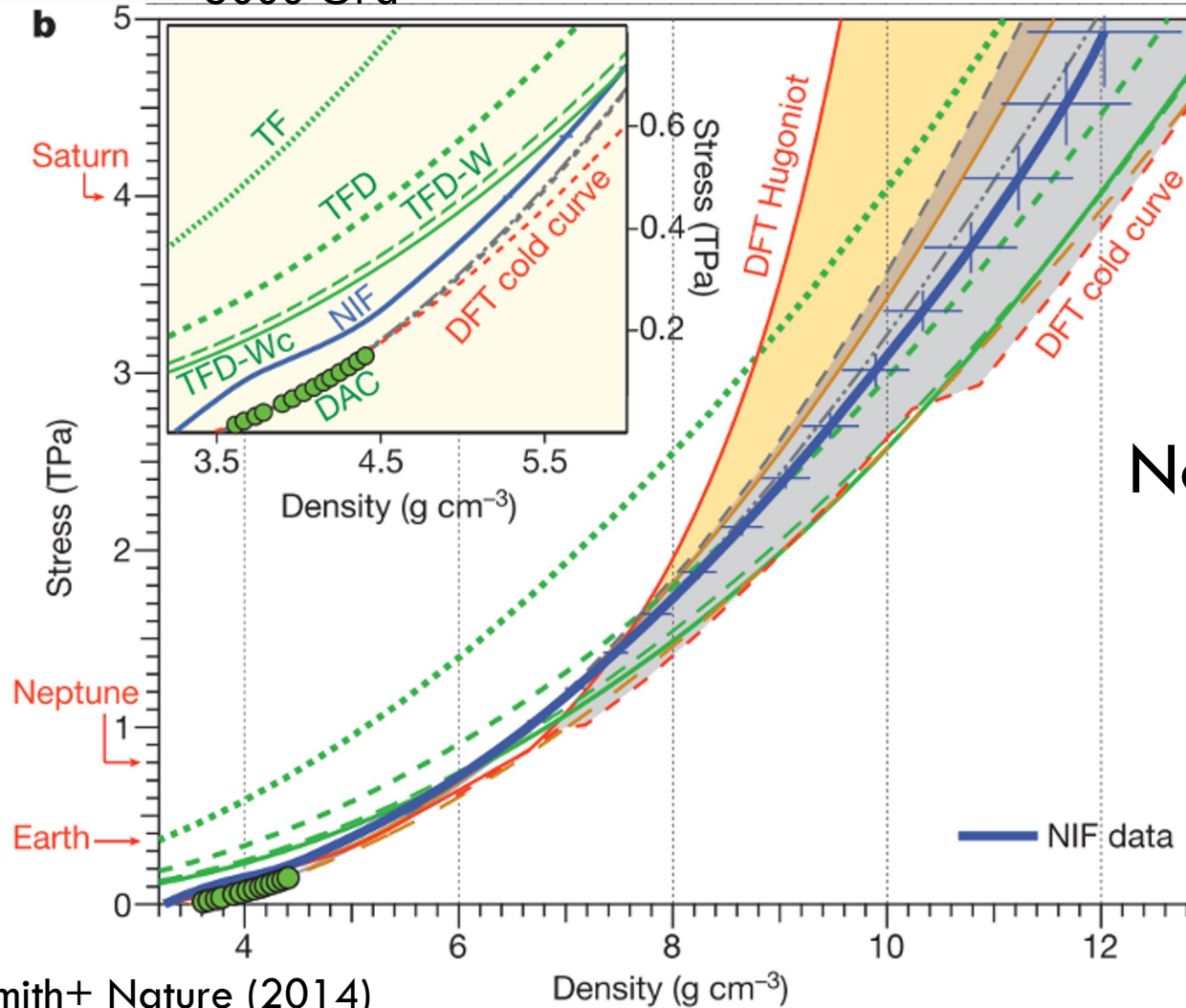
Gonzalez+, PRB (2021)

Diamond



from ab initio simulations

2000 GPa



No signature of bc8

Article

Metastability of diamond ramp-compressed to 2 terapascals

Lazicki+, Nature (2021)

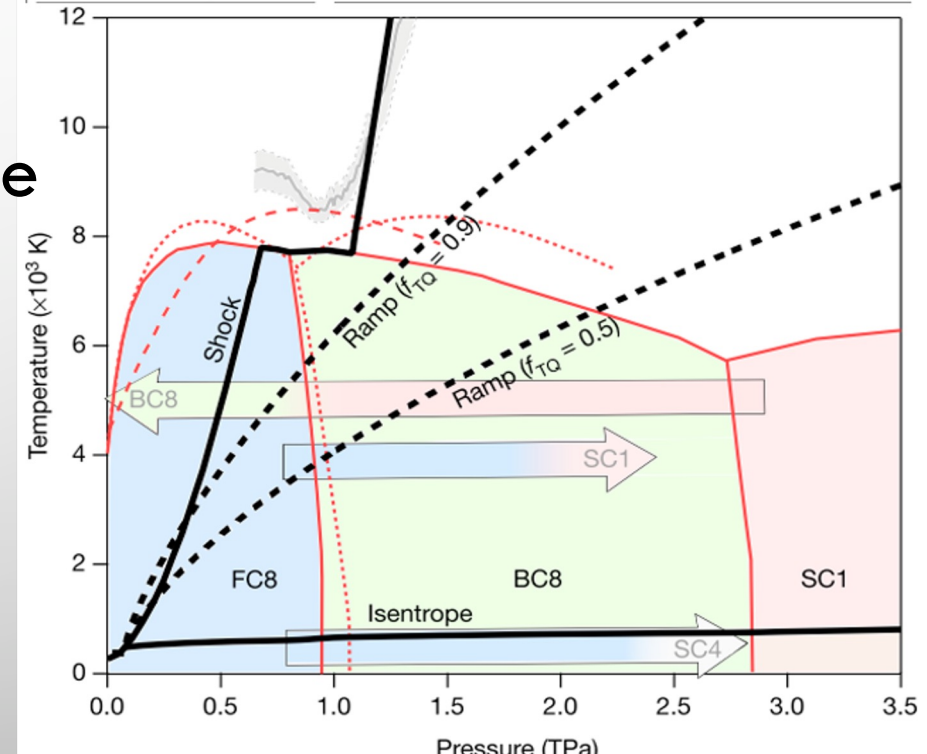
<https://doi.org/10.1038/s41586-020-03140-4>

Received: 1 June 2020

Accepted: 26 October 2020

Published online: 27 January 2021

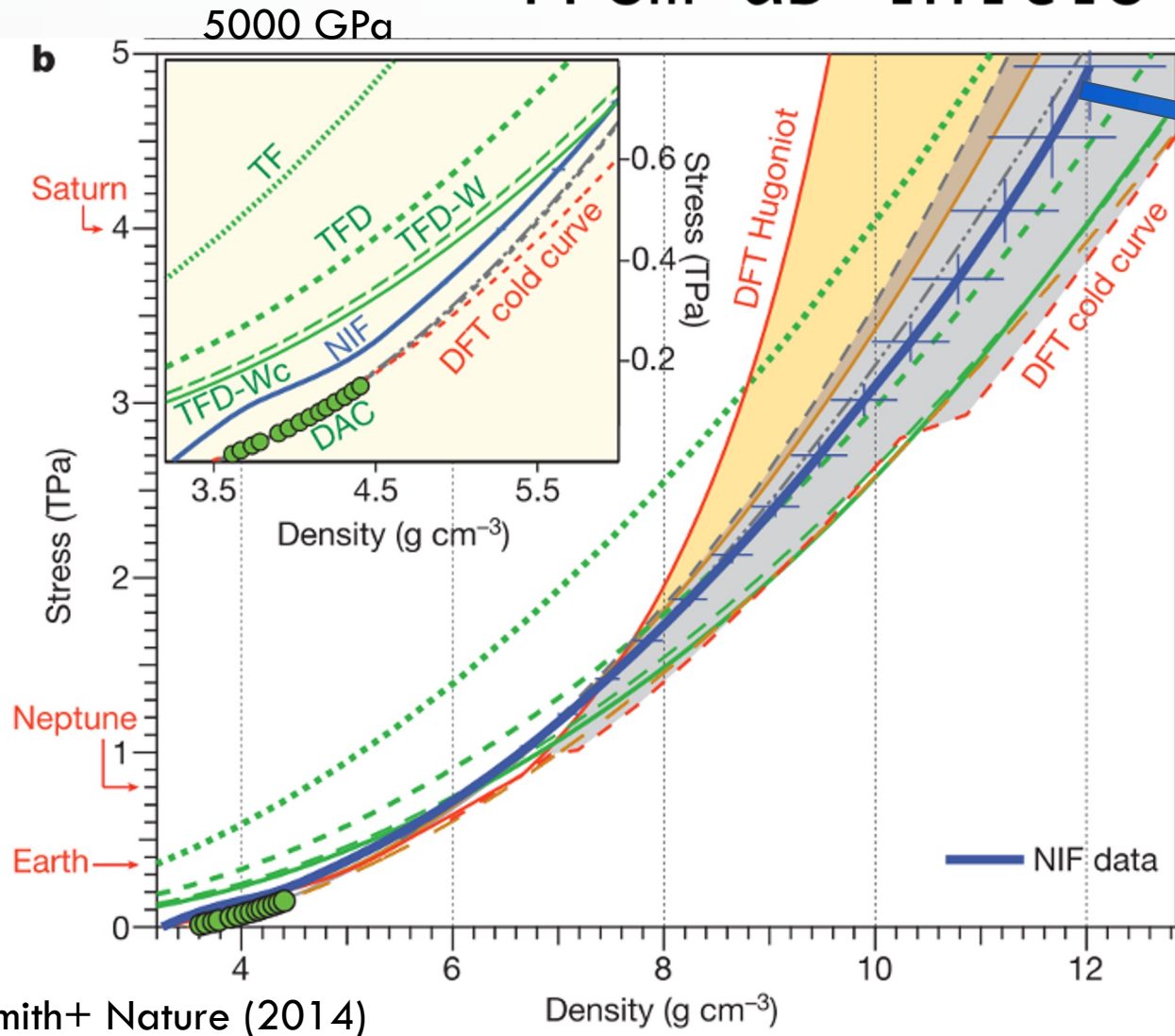
A. Lazicki^{1,2*}, D. McGonegle³, J. R. Rygg^{3,4,5}, D. G. Braun¹, D. C. Swift¹, M. G. Gorman¹, R. F. Smith¹, P. G. Helghway², A. Higginbotham⁶, M. J. Suggit⁷, D. E. Fratanduono¹, F. Coppari¹, C. E. Wehrenberg¹, R. G. Kraus¹, D. Erskine¹, J. V. Bernier¹, J. M. McNaney¹, R. E. Rudd¹, G. W. Collins^{2,4,5}, J. H. Eggert¹ & J. S. Wark²



Ramp compression model

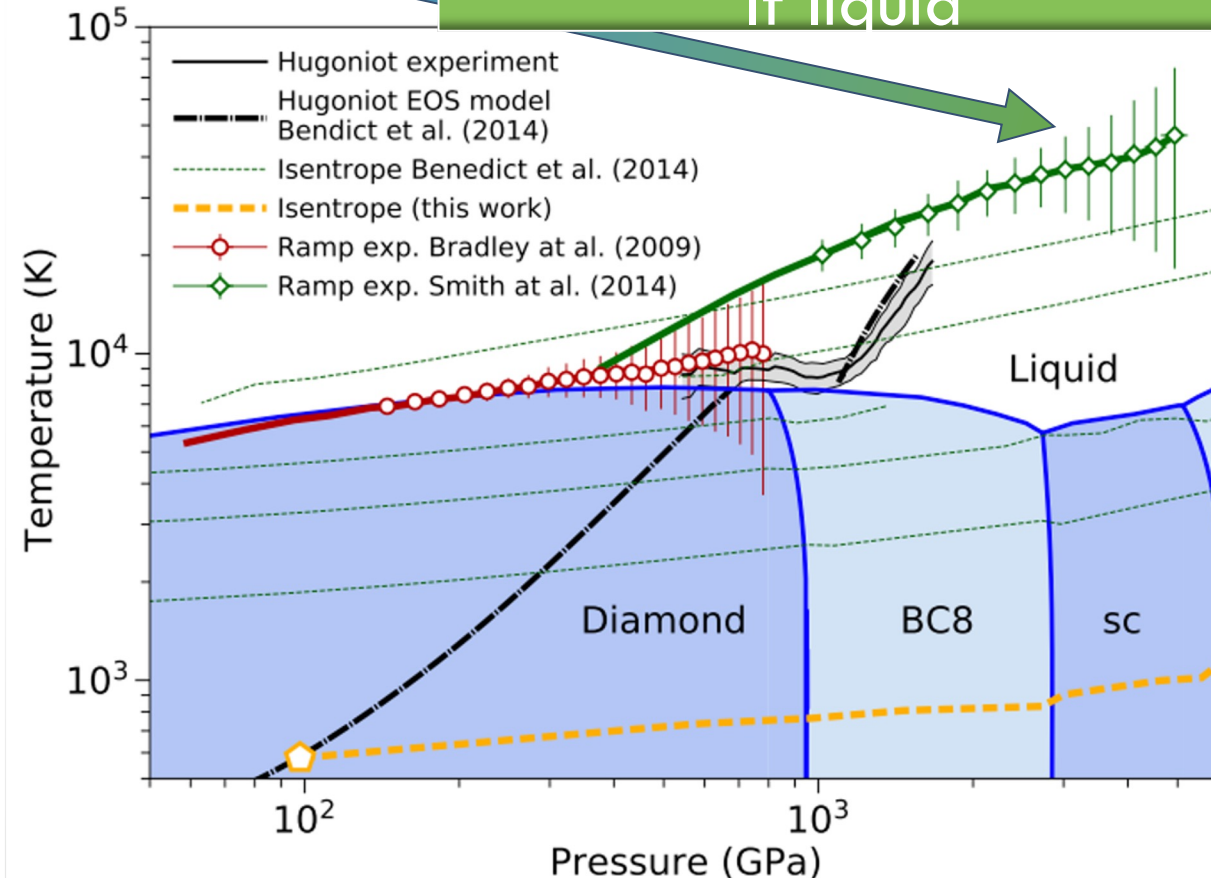
Gonzalez+, PRB (2021)

from ab initio simulations



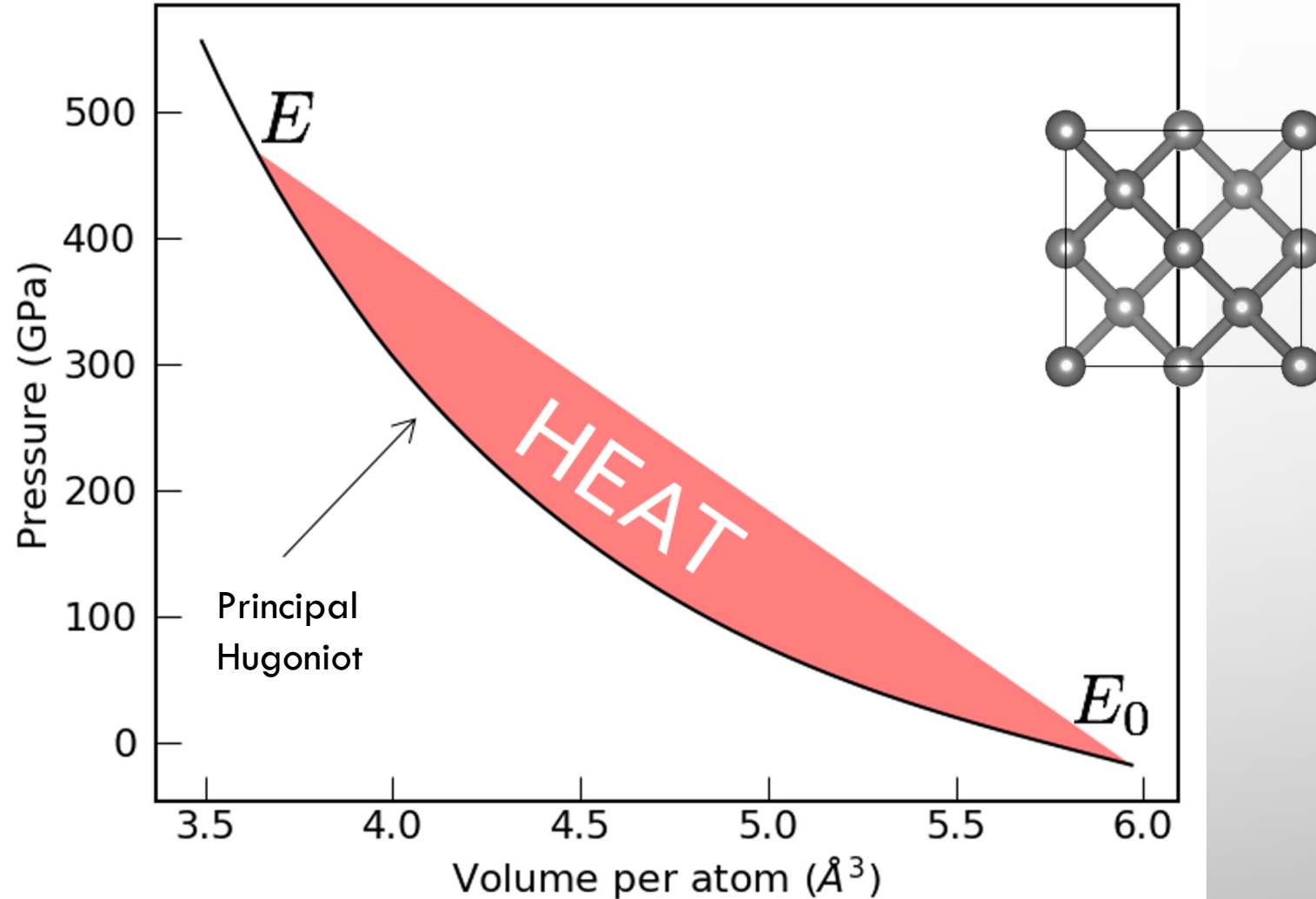
$T > 20\,000\text{ K}$

If liquid



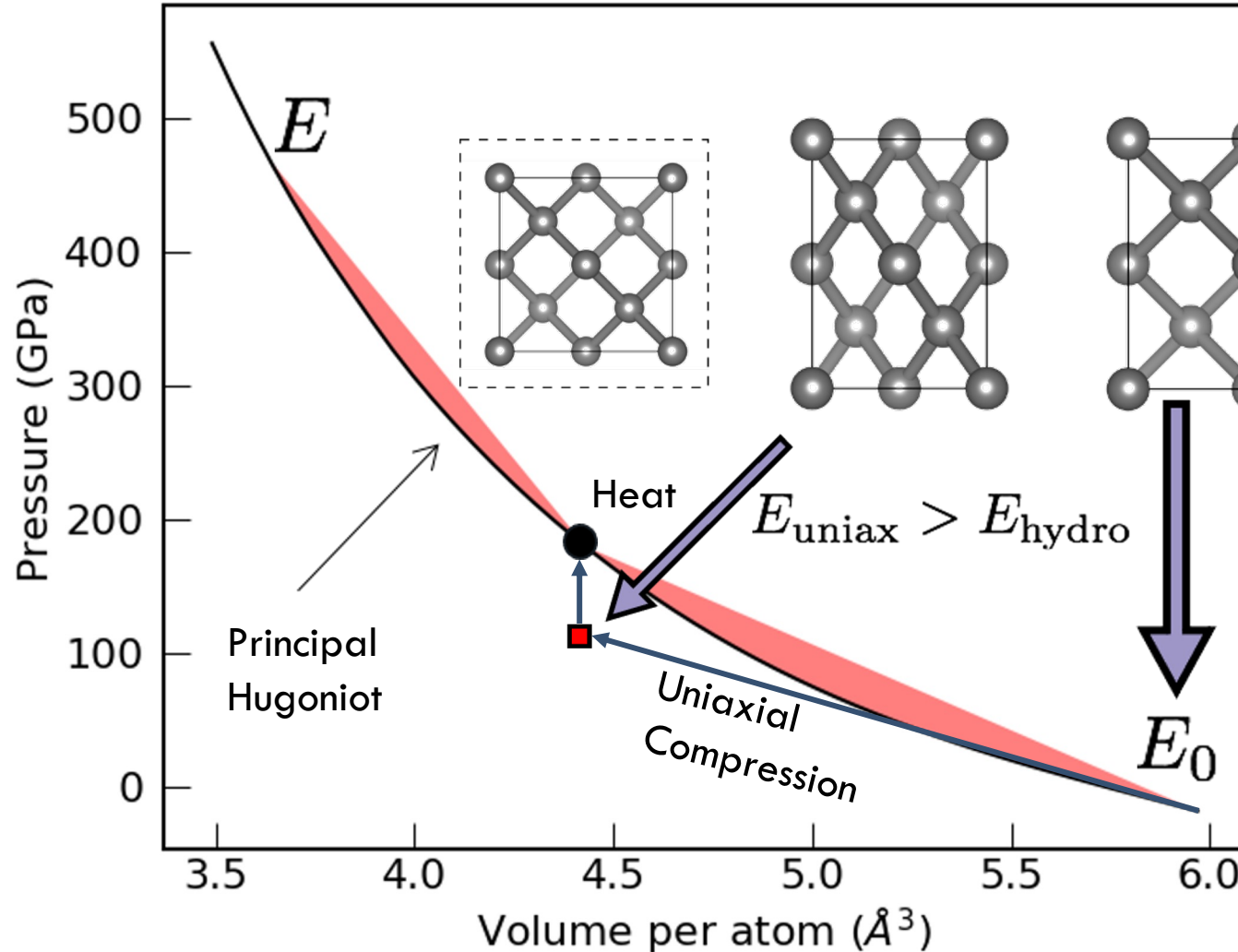
Ramp \sim multishocks

$$E - E_0 = \frac{1}{2}(P + P_0)(V_0 - V)$$

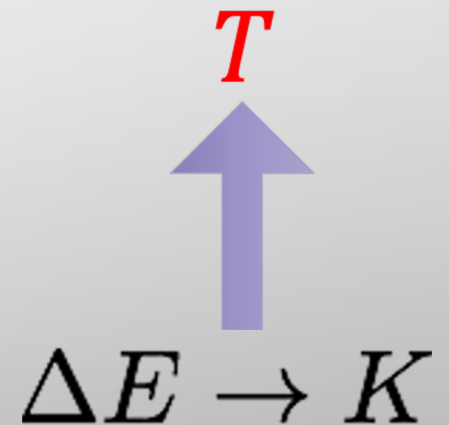


Ramp \sim multishocks

$$E - E_0 = \frac{1}{2}(P + P_0)(V_0 - V)$$

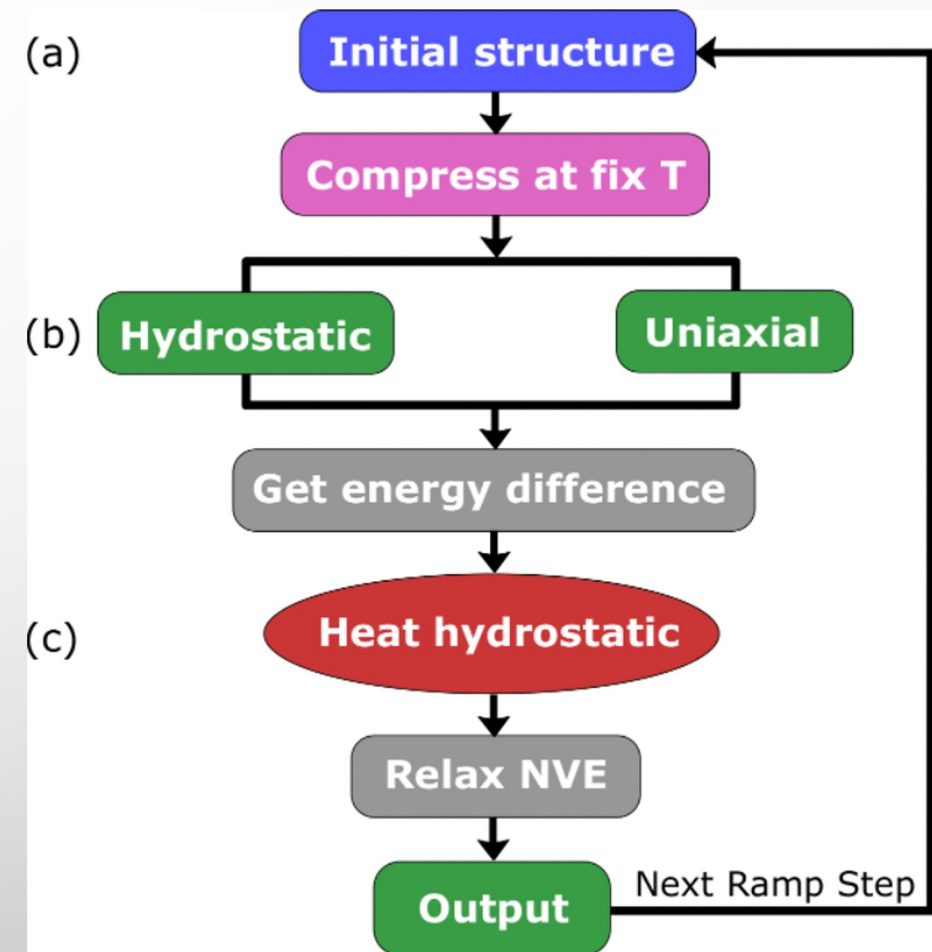
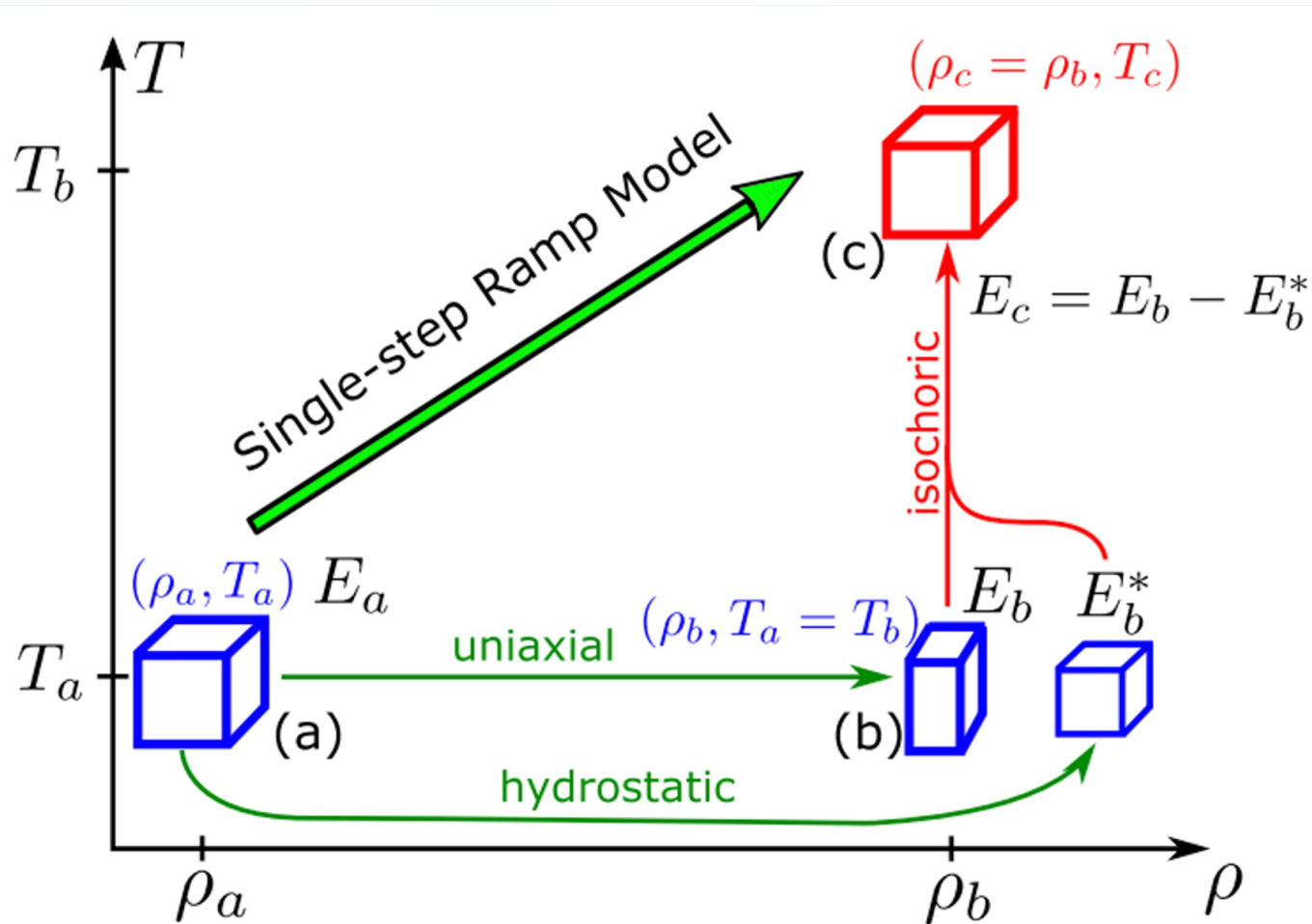


1. Initial state
2. Uniaxial compression
3. Ediff \rightarrow Kinetic energy

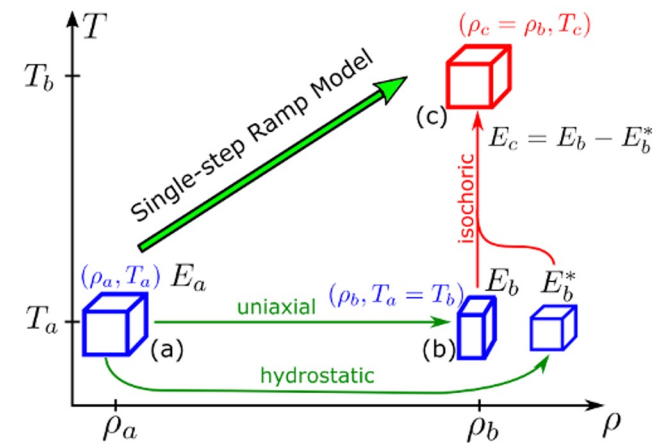
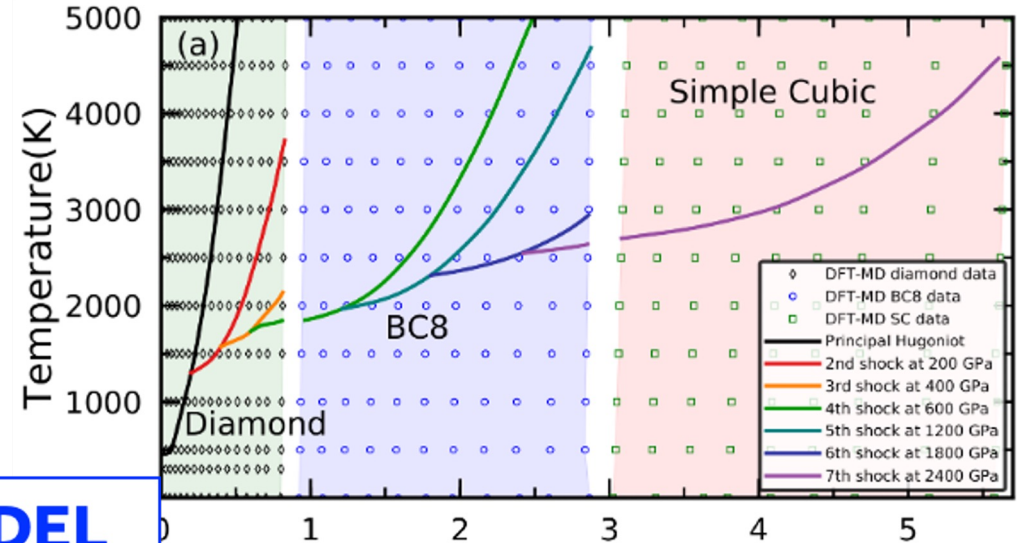
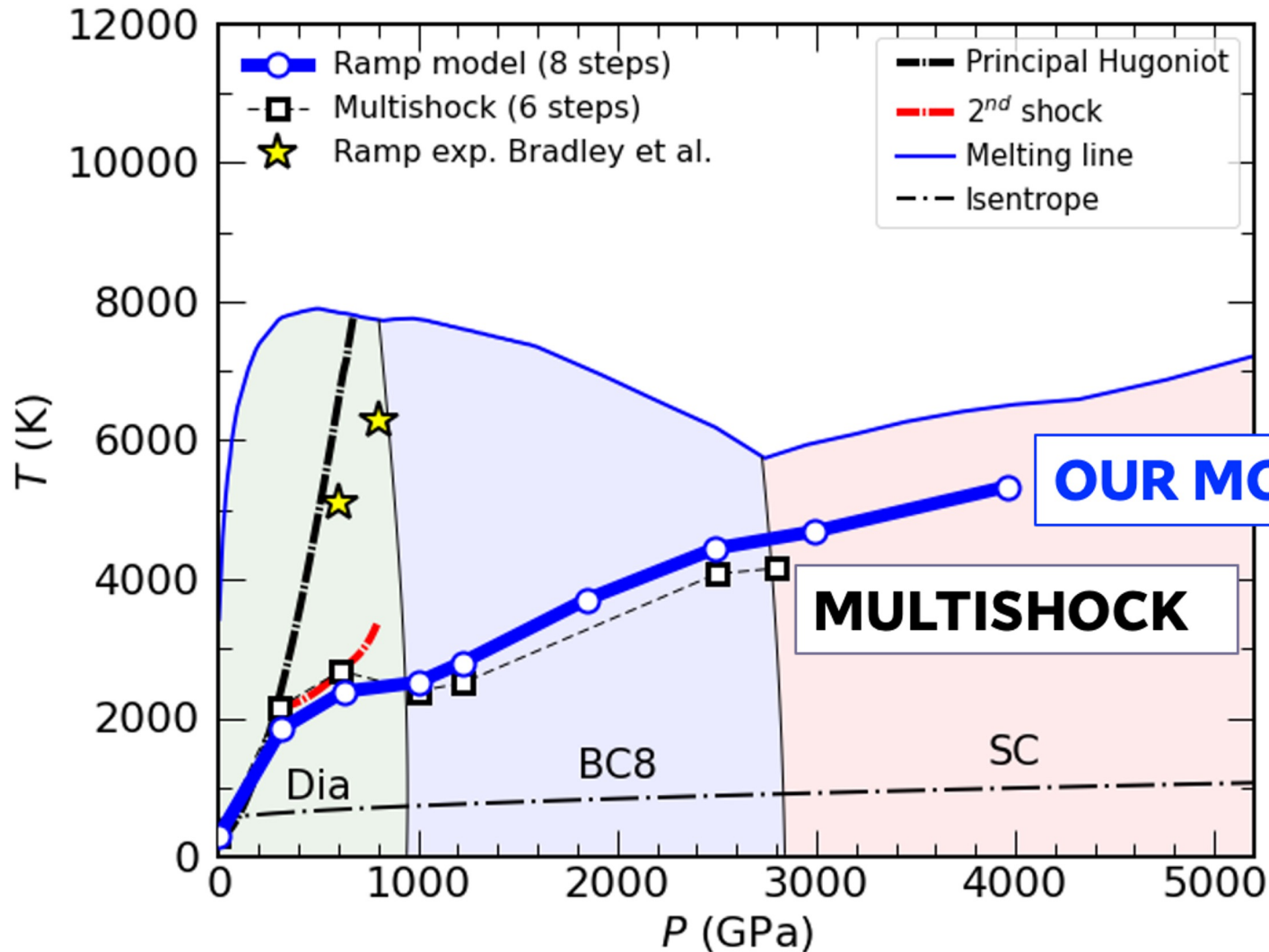


OUR MODEL OF RAMP COMPRESSION

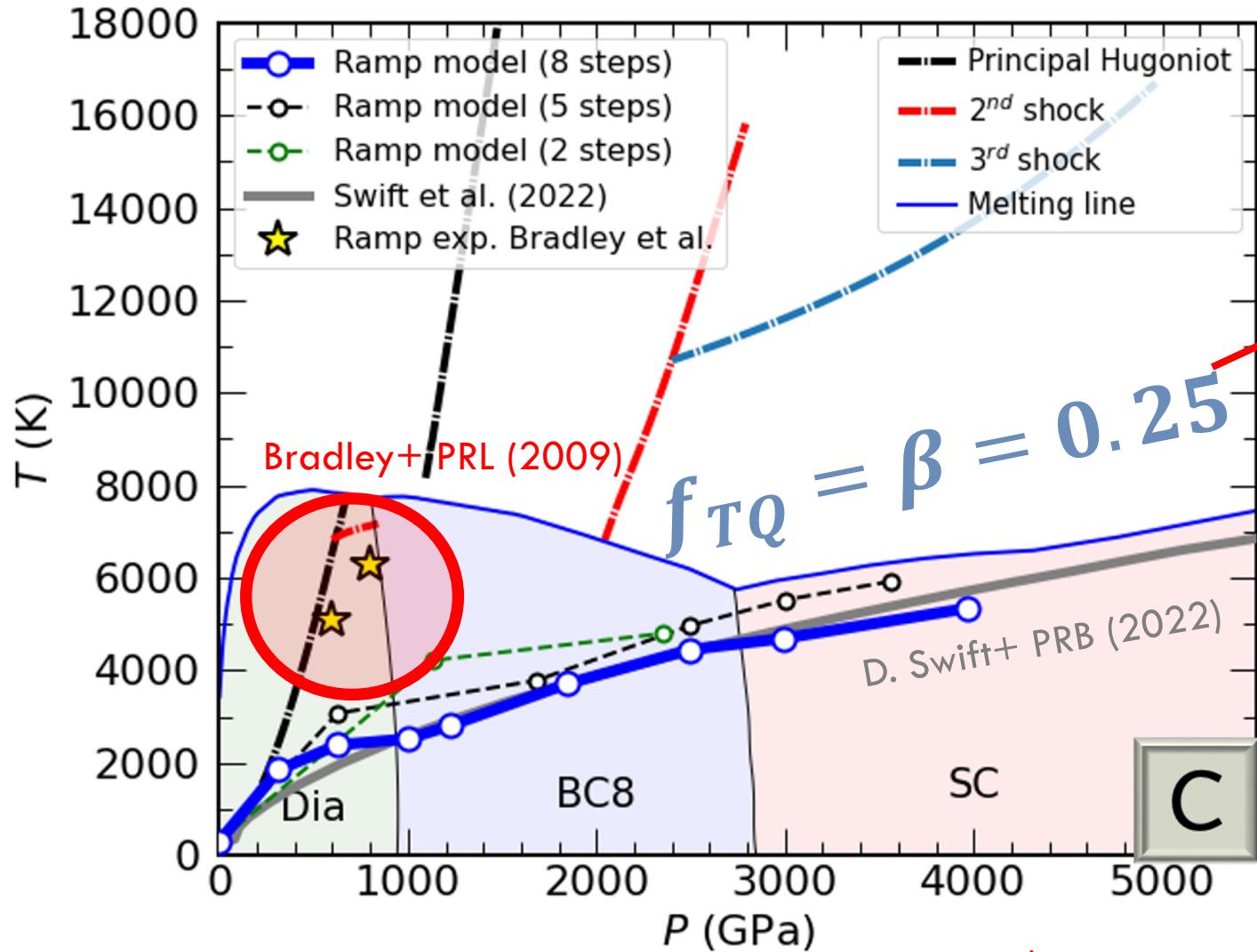
for ramp compression from ab initio simulations



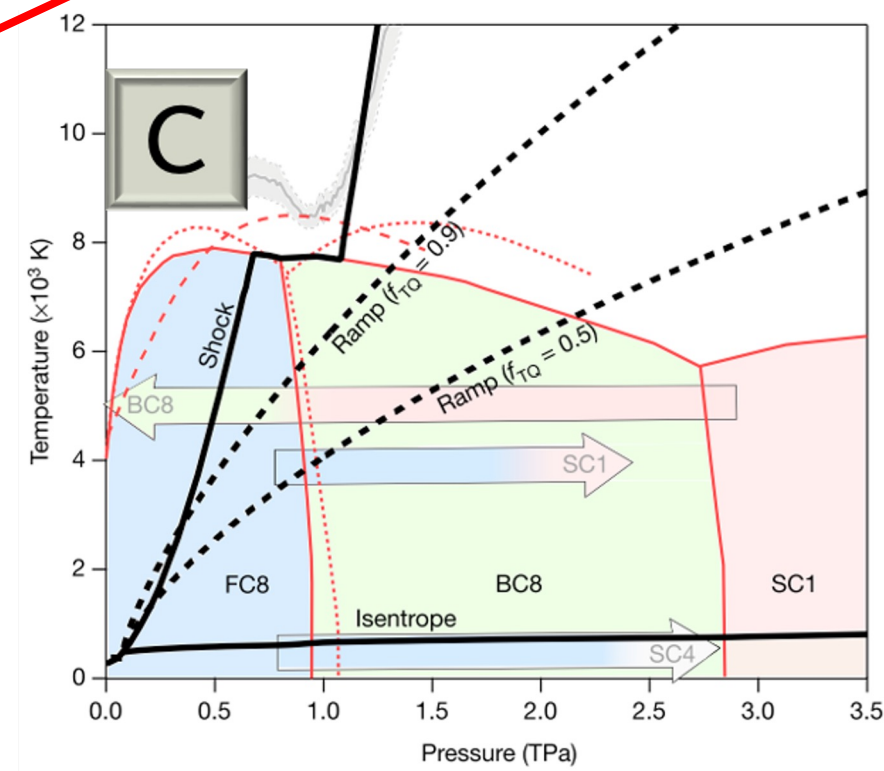
MULTISHOCKS / OUR MODEL



MULTISHOCKS / OUR MODEL



$$P_{isen} = \sigma_x - \frac{2}{3} Y - \int \beta dW_p - \gamma \rho_H (E_H - E_{isen})$$



* Diamond is weaker than expected
 * 3/4 plastic work absorbed by defects!

Lazicki+, Nature (2021)

OUTLINE

1. Planetary Interiors

- A diluted core in Jupiter
- Rock/Ice mixtures in water planets

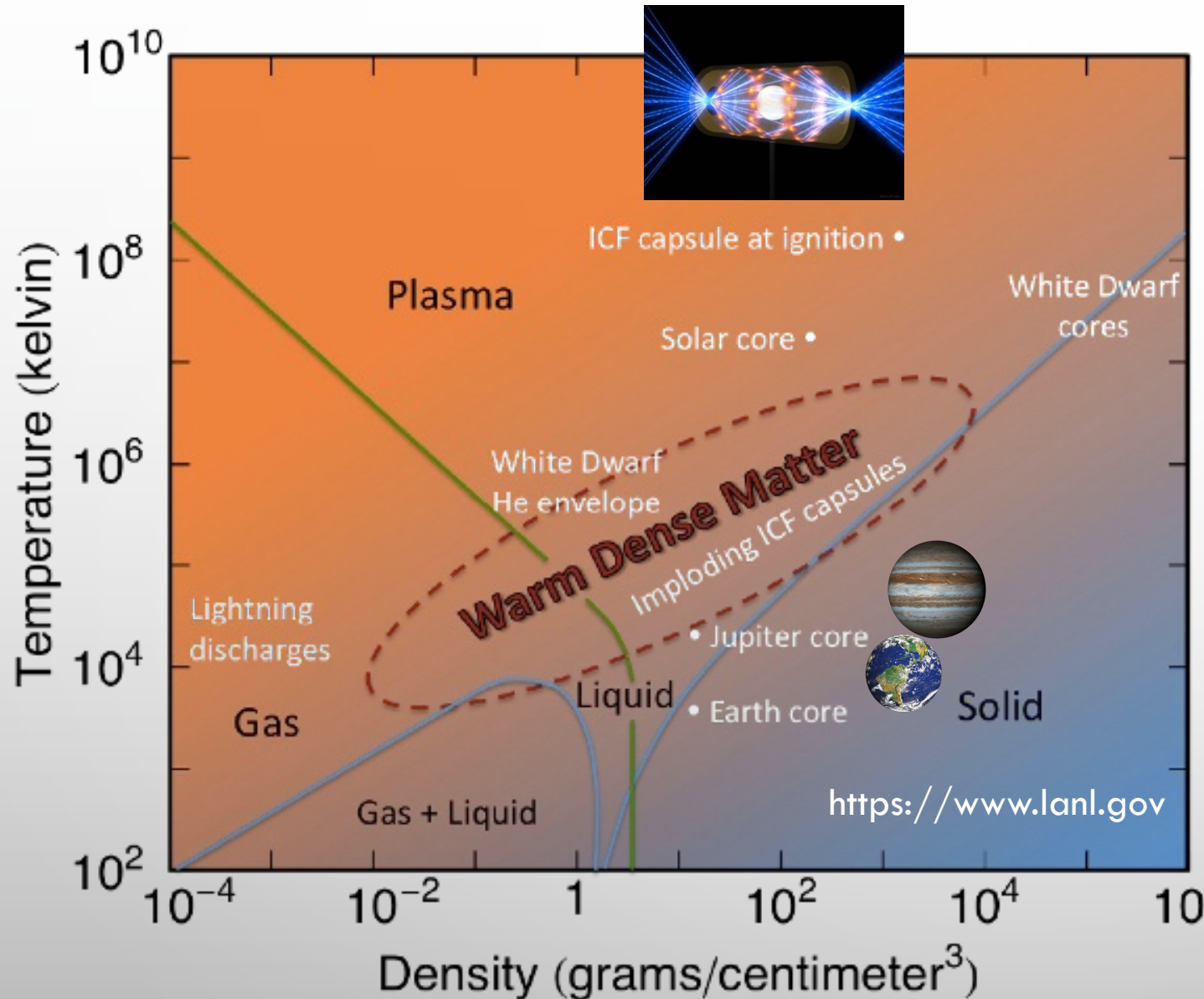
2. High Pressure Phase Transitions

- Be & MgO: melting and anharmonicities
- Melting of SiO_2
- Ramp compression from DFT

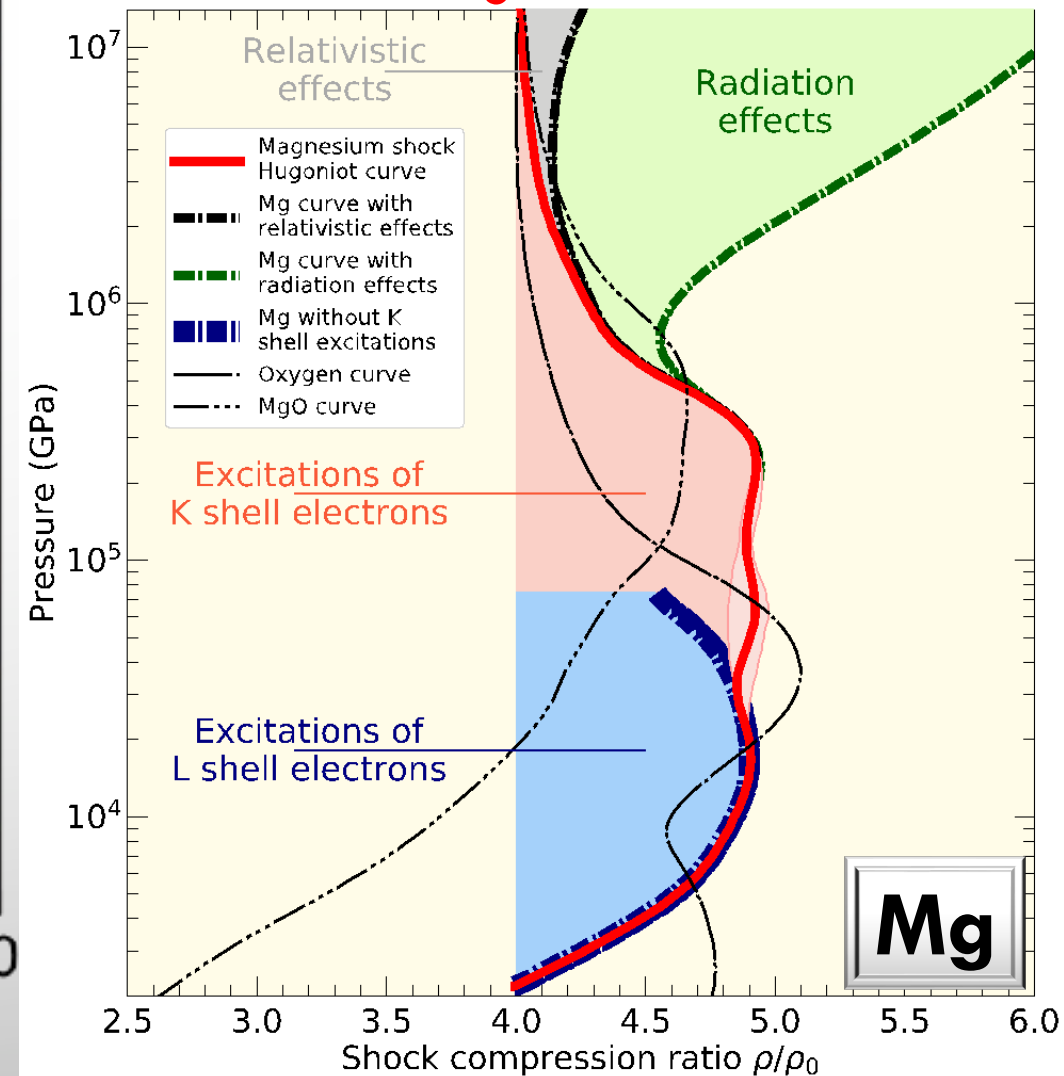
3. Warm Dense Matter

- Warm dense silicates: Mg, MgO & MgSiO_3
- FPEOS

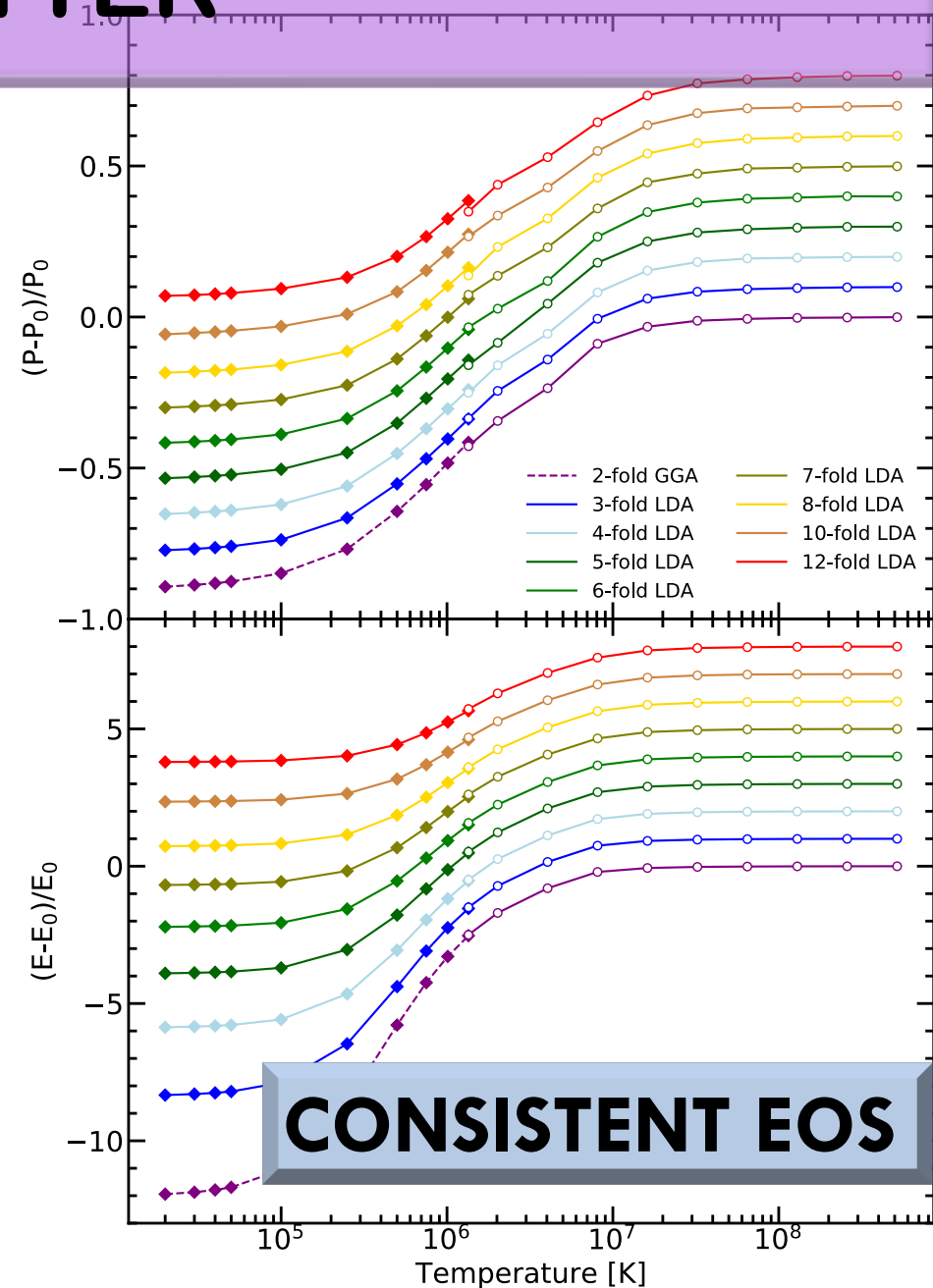
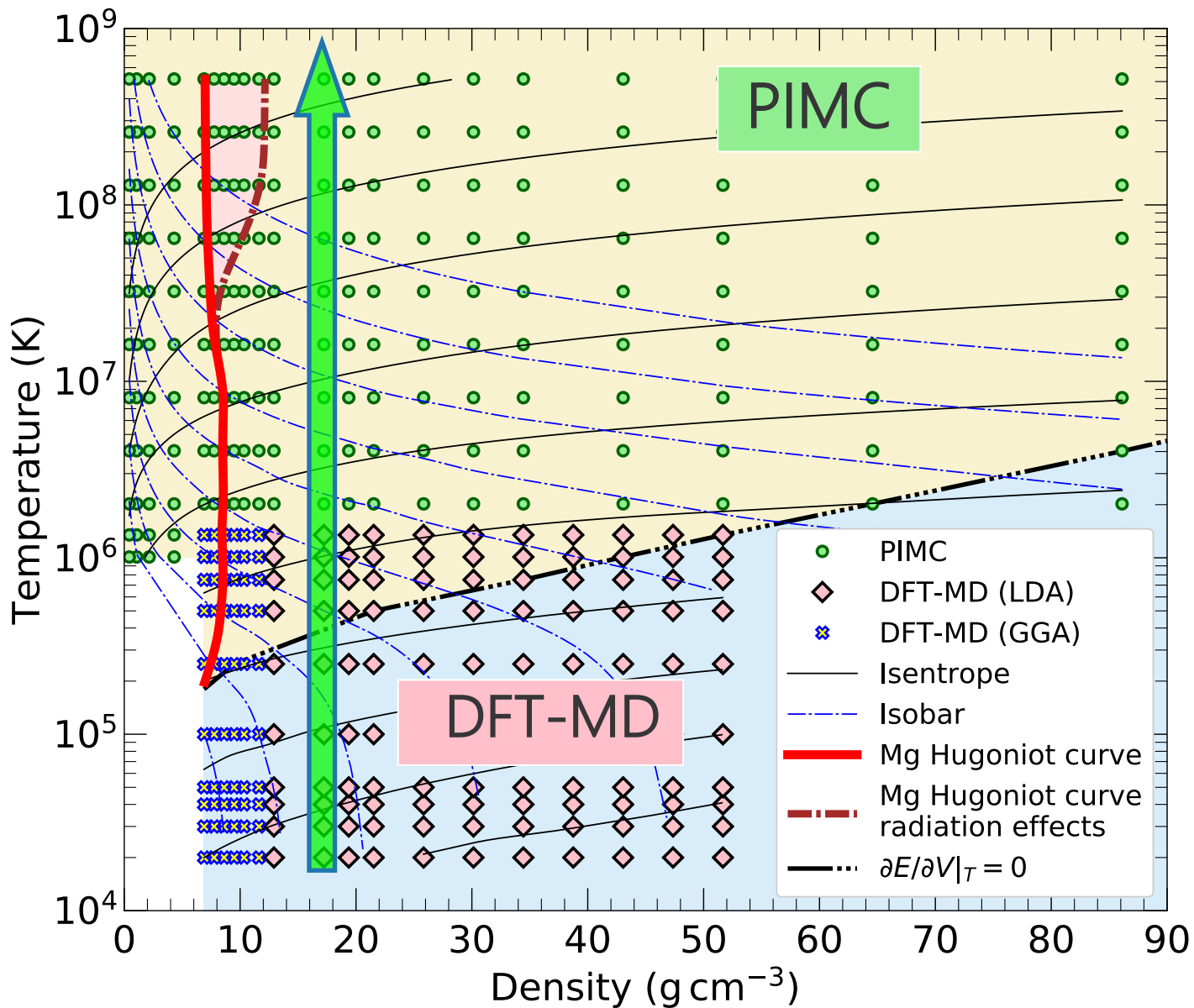
WARM DENSE MATTER



Shock compression: The Hugoniot curve

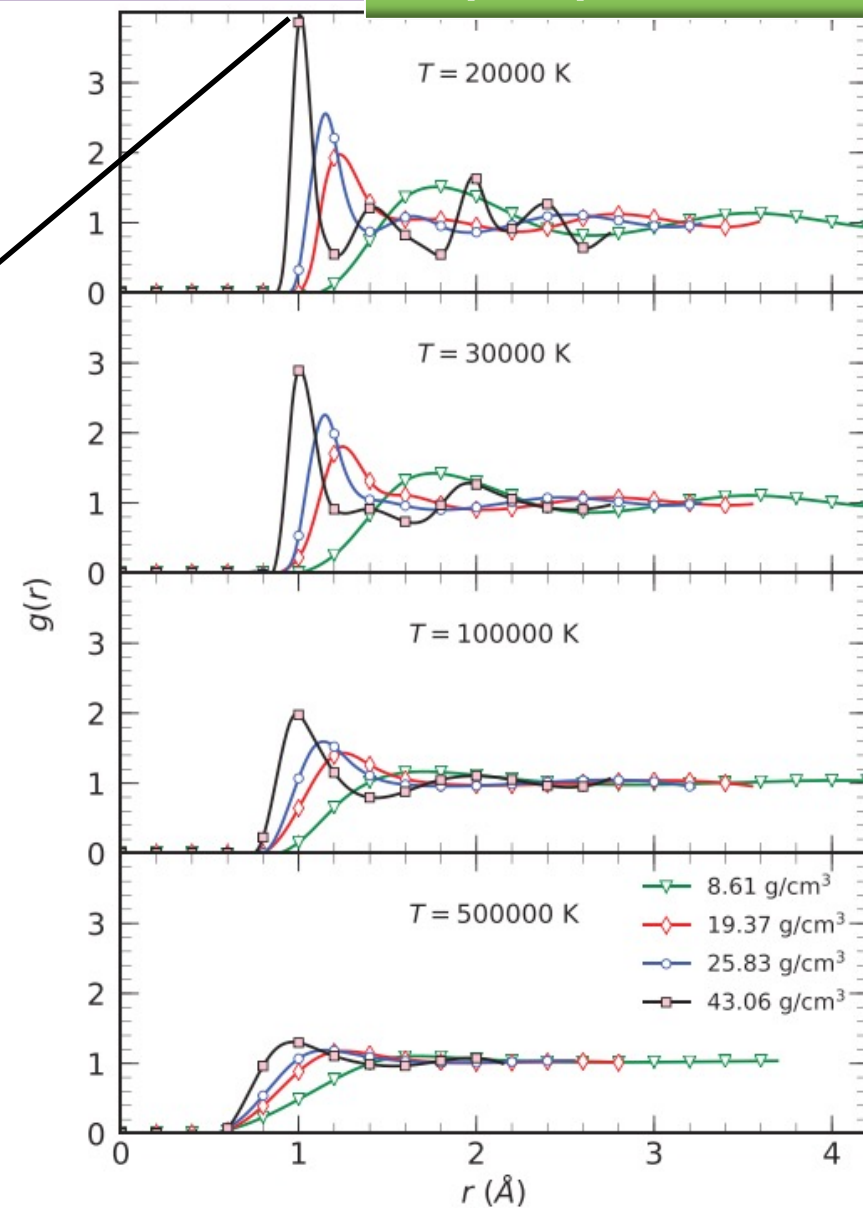
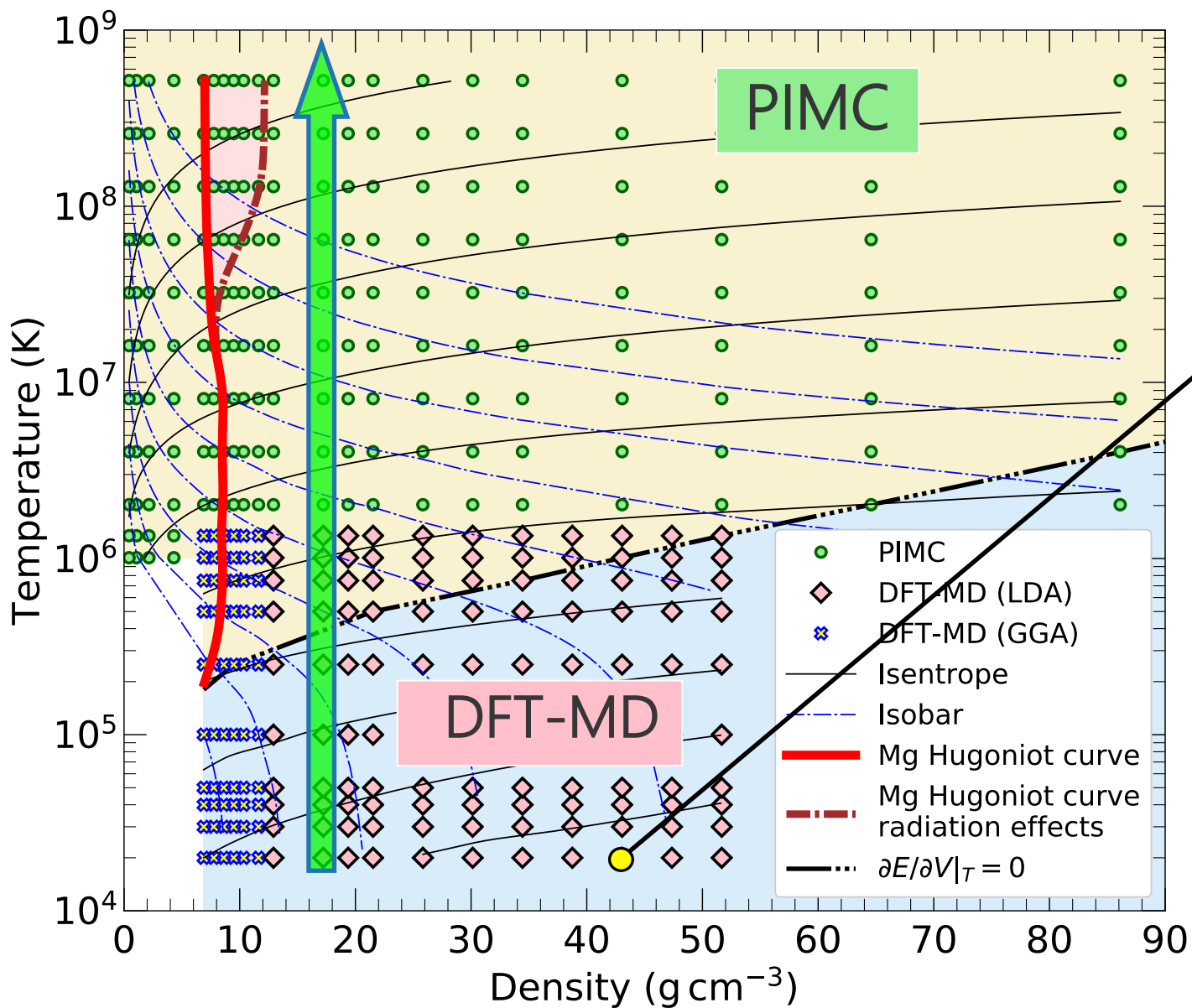


WARM DENSE MATTER

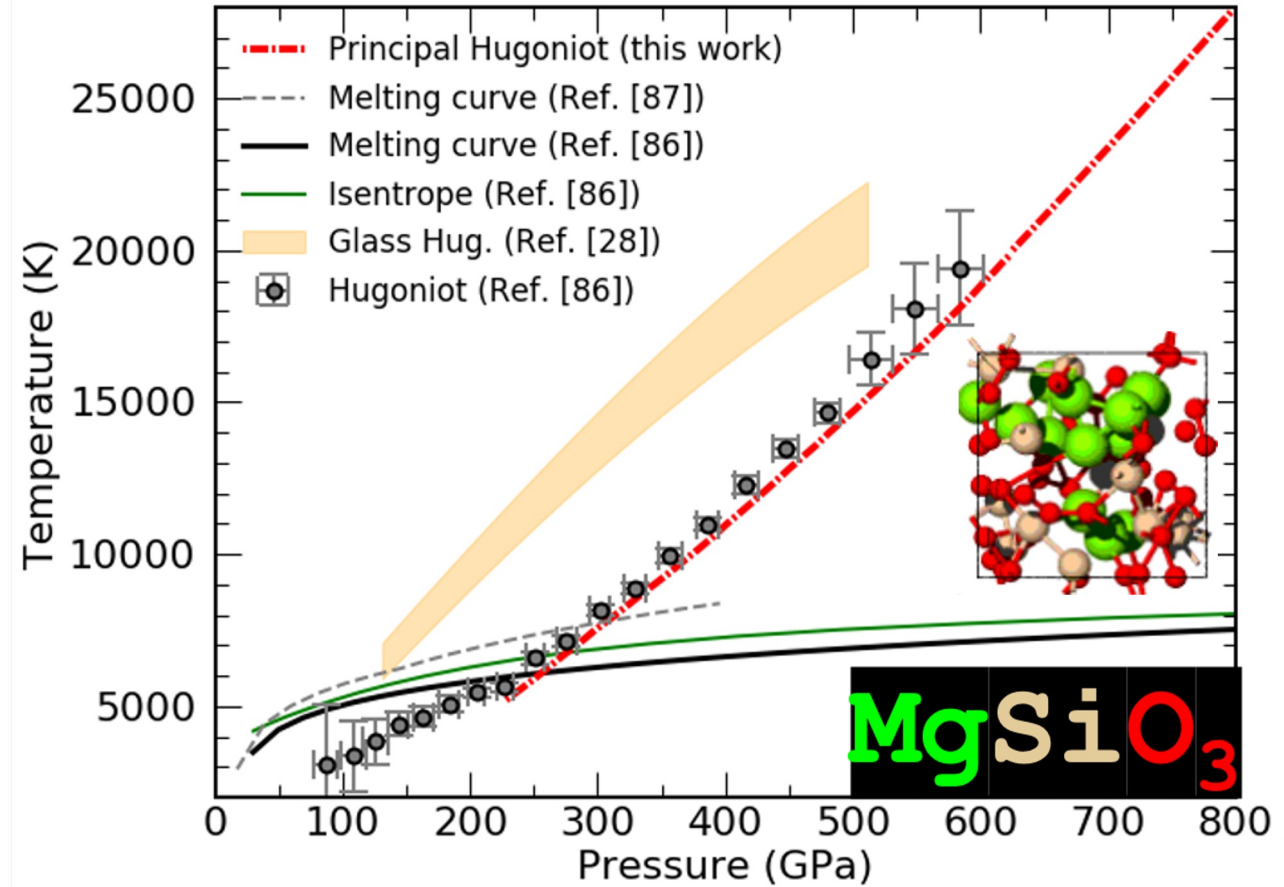
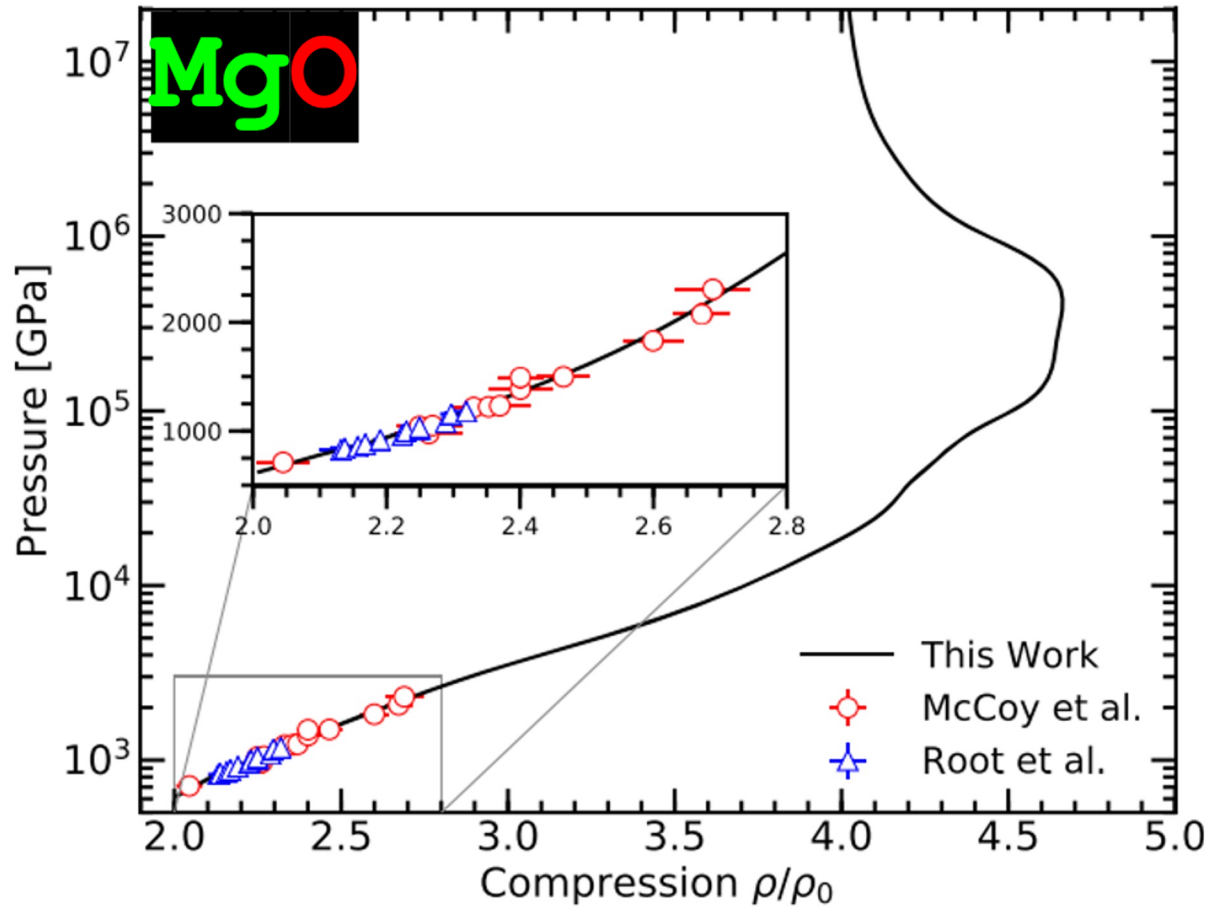


WARM DENSE MATTER

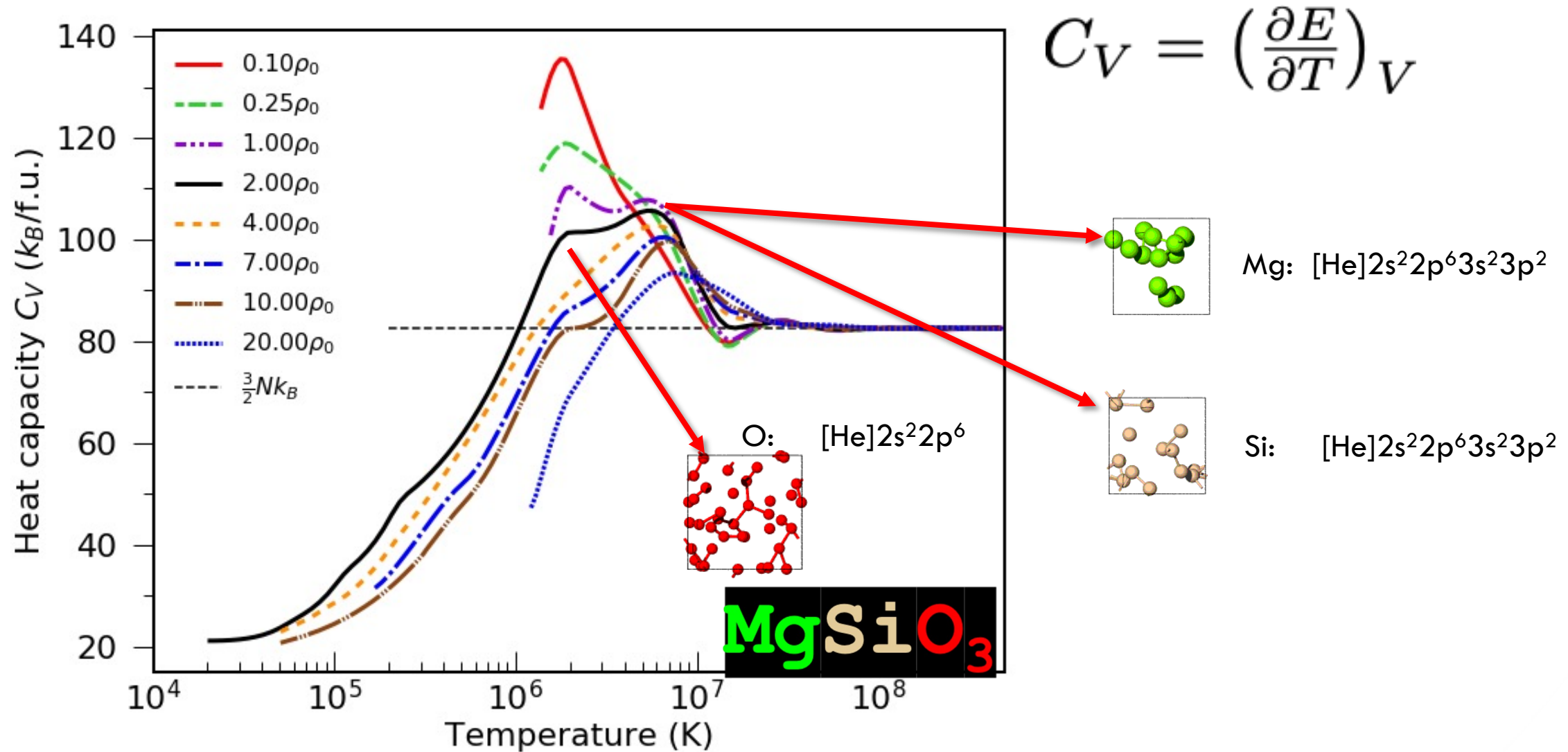
$g(r)$ Structural properties



WARM DENSE MATTER

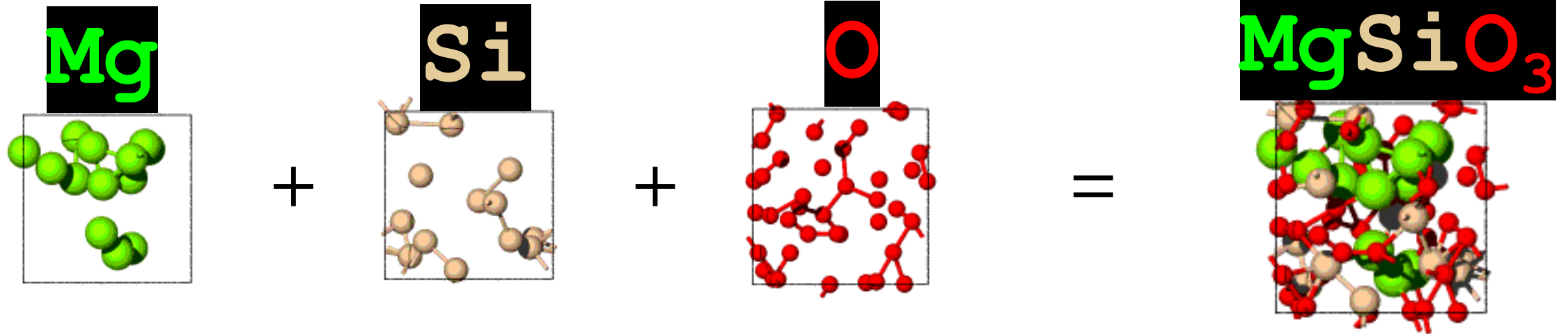


WARM DENSE MATTER



WARM DENSE MATTER

Non ideal mixing



Linear Mixing Approximation: (at constant P and T)

$$V_{mix} = N_1 V_1 + N_2 V_2 + N_3 V_3$$

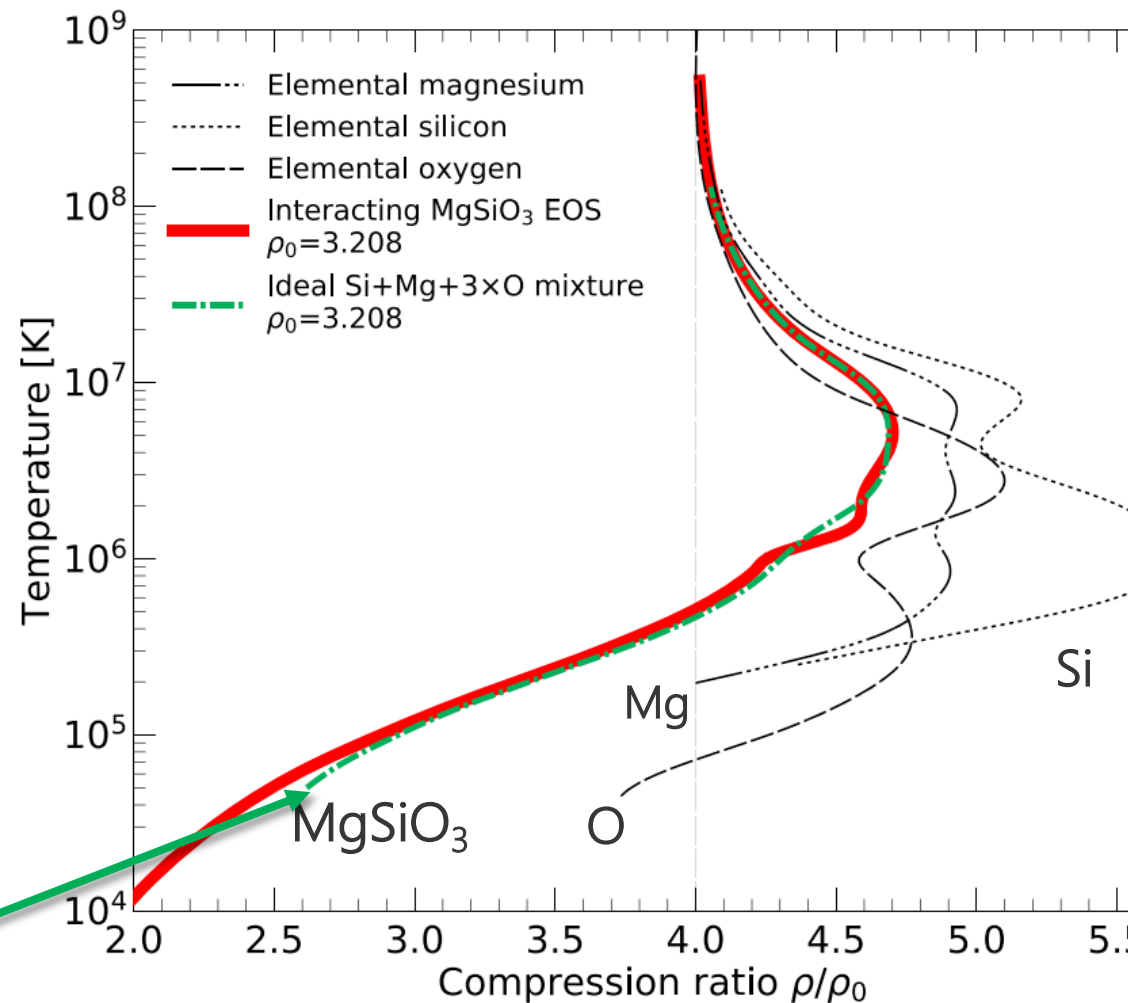
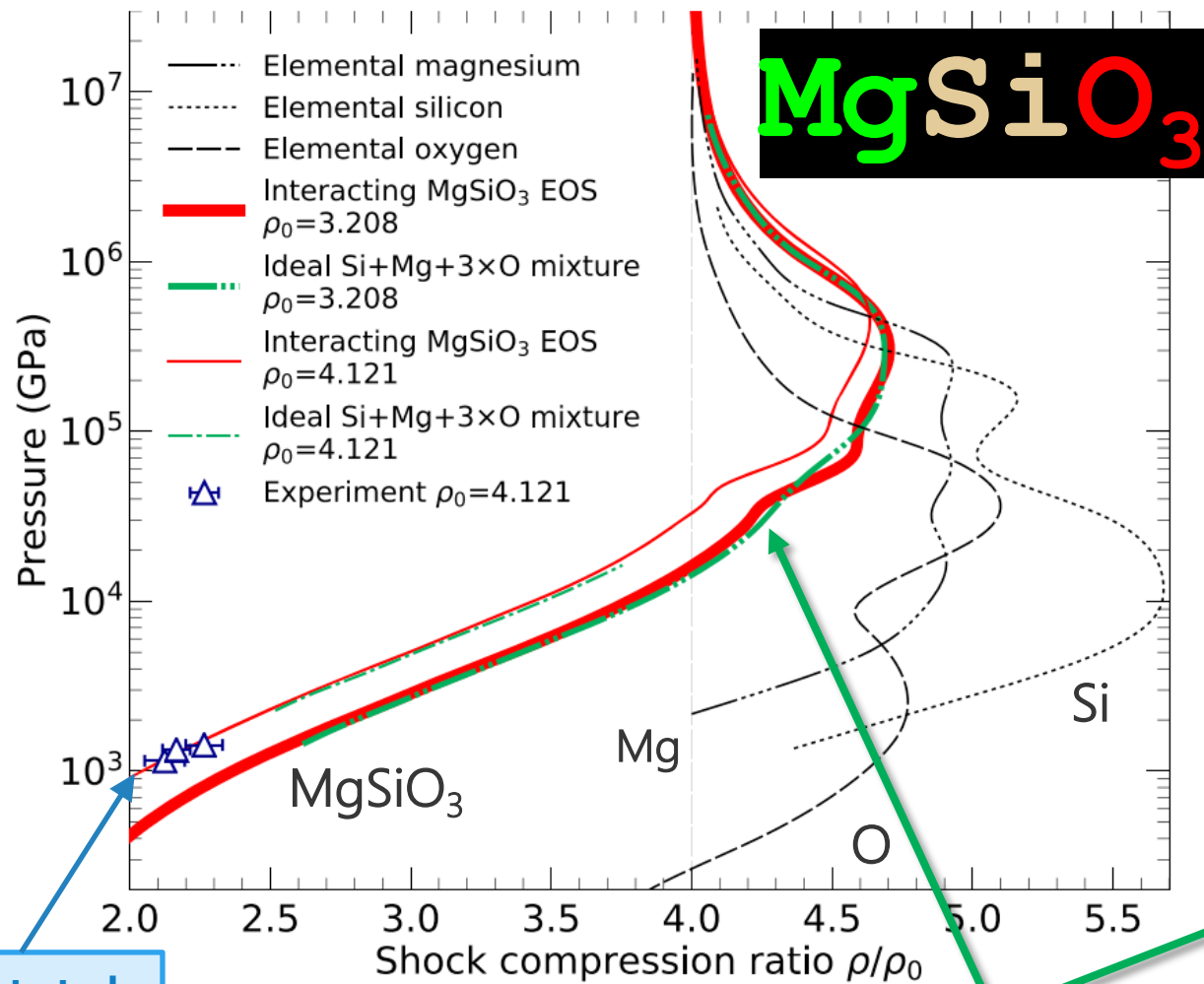
$$m_{mix} = N_1 m_1 + N_2 m_2 + N_3 m_3$$

$$E_{mix} = N_1 E_1 + N_2 E_2 + N_3 E_3$$

(additive volume rule)

$$\rho_{mix} = m_{mix} / V_{mix}$$

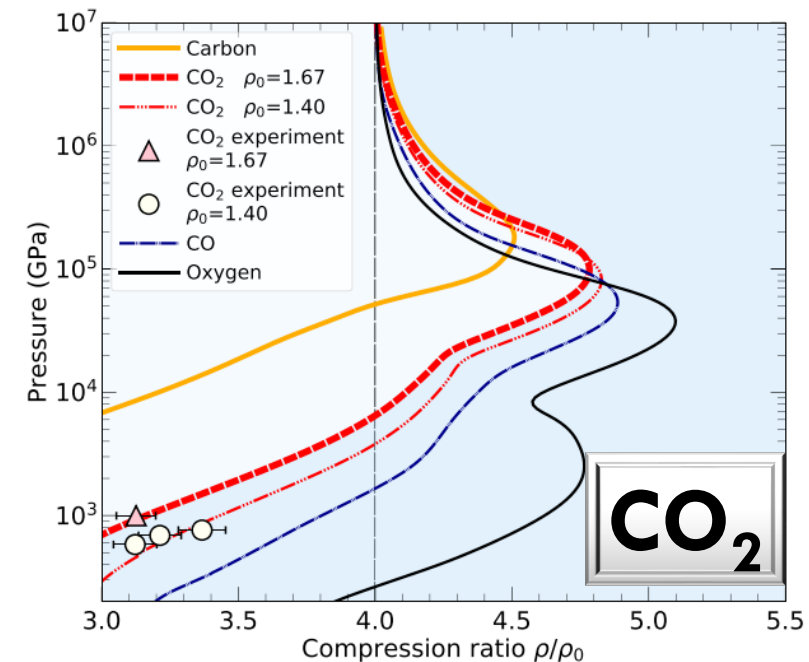
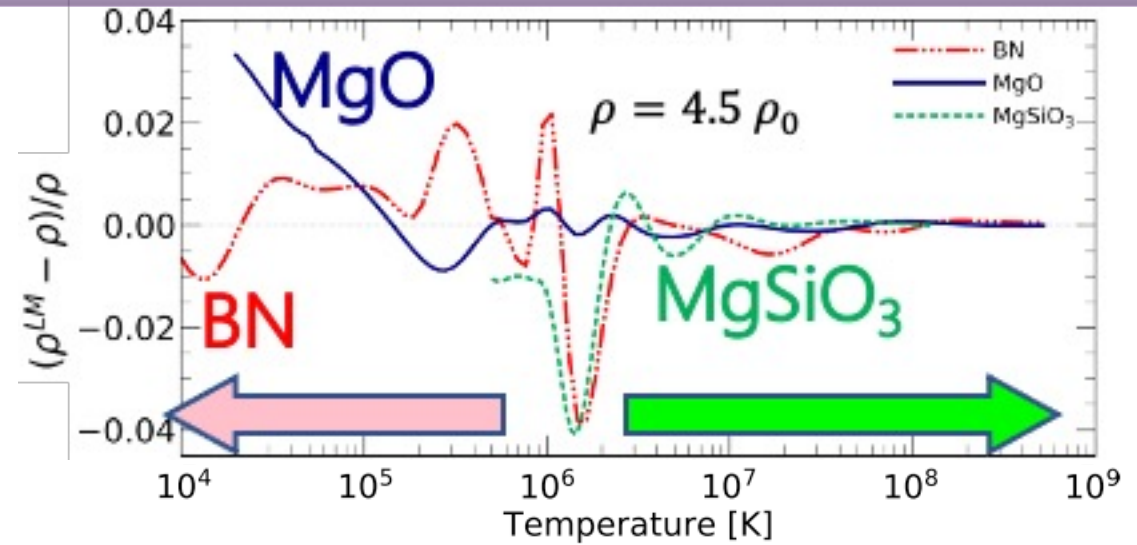
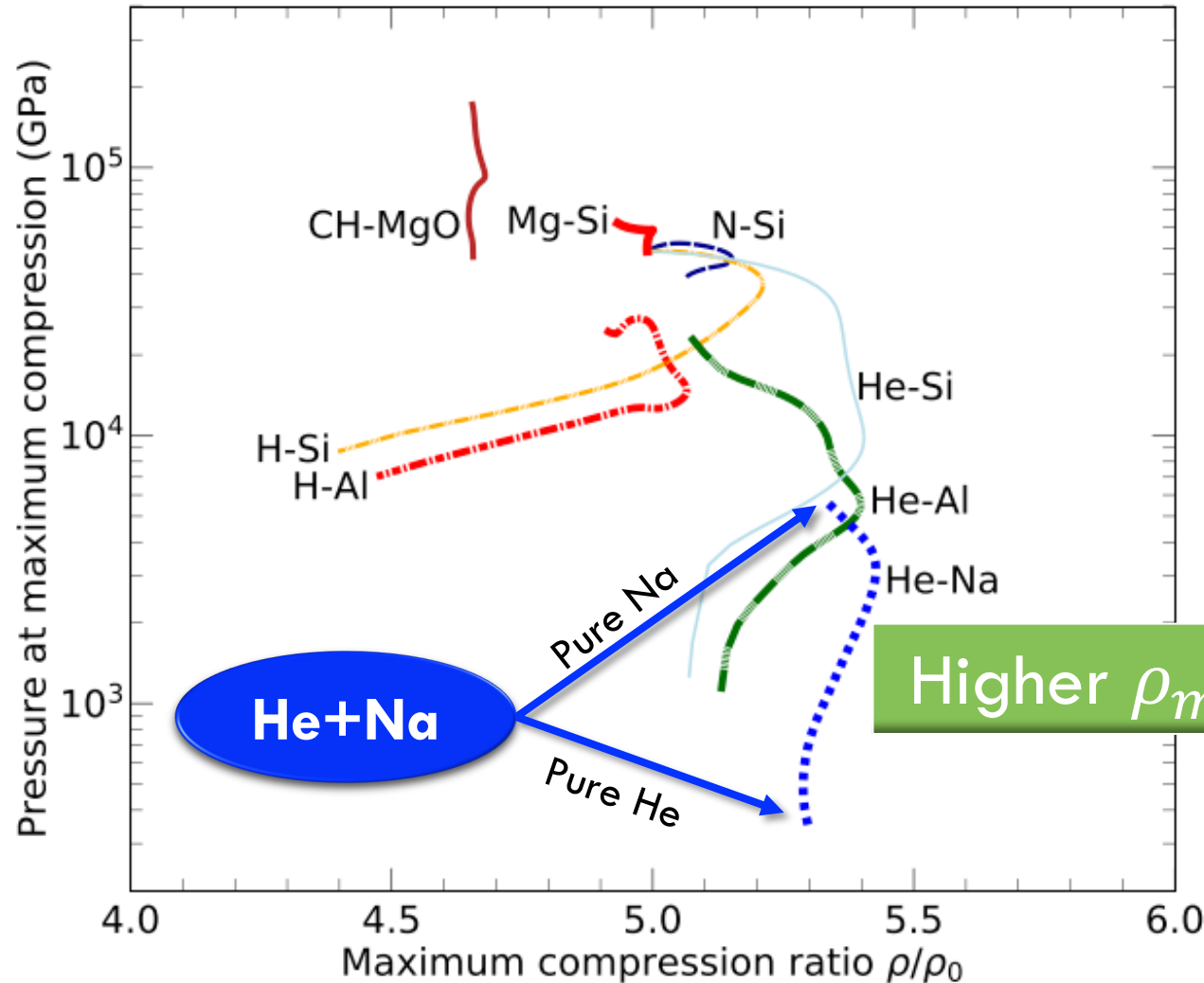
WARM DENSE MATTER



Ideal mixing

FIRST PRINCIPLES EQUATION OF STATE (FPEOS)

<http://militzer.berkeley.edu/FPEOS>



CONCLUSIONS

1. Planetary interiors: mixing, erosion, crystallization.
2. Thermodynamic Integration: melting, anharmonicity
3. Z method agrees with two-phase and ΔG .
4. Ramp compression: better models needs
5. Validated Linear Mixing for MgO , MgSiO_3 , and BN plasmas.
6. <http://militzer.berkeley.edu/FPEOS/>

Thanks!

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